

# 湘中龙山大型金锑矿床成矿时代研究 ——黄铁矿 Re-Os 和锆石 U-Th/He 定年\*

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**Abstract** The Longshan deposit is one of the most important Au-Sb deposits in central Hunan Province. However, the mineralization age of this deposit is still not well constrained due to lacking of suitable minerals for conventional radiometric dating, limiting the recognition of its genesis. Pyrite Re-Os isotopic dating has been proved to be a reliable technique to determine the age of hydrothermal sulfides with low contents of Re and Os, which can be thus used to constrain the age of epithermal mineralization. The zircon U-Th/He isotopic dating is a newly developed technique, which can provide a potential constraint on the ages of low-temperature metallogenic events. Re-Os isotopic dating of the hydrothermal pyrites and U-Th/He isotopic dating of the zircons recording metallogenic event were carried out in this paper, aiming to constrain the mineralization age of the Longshan Au-Sb deposit. Dating results show that the Re-Os isochron age of hydrothermal pyrites is  $195 \pm 36$  Ma, corresponding to Late Indosinian, whereas the U-Th/He ages of zircons from altered wallrock are 51.2 ~ 133.3 Ma. After correction by Ft, the U-Th/He ages are distributing in the range of 93.78 to 258.29 Ma with an average of  $160.7 \pm 7.3$  Ma, corresponding to Early Yanshanian. It can be thus concluded that the Longshan Au-Sb deposit may have experienced two episodes of mineralization (i. e., ~200 Ma and ~160 Ma) or the Sb-Au mineralization may have occurred at ~200 Ma which was remobilized by magmatic event at ~160 Ma. Pyrite Re-Os age represented the timing of Au-Sb mineralization, whereas the zircon U-Th/He age represented the timing of the second thermal event. Thus, whether the Longshan Au-Sb deposit formed due to the single mineralization event at ~200 Ma, or also experienced an overlapping of the mineralization event at ~160 Ma, the two ages were roughly consistent with those of two magmatic events in the district, indicating that magmatism may have played a vital role in the formation of the Longshan Sb-Au deposit.

**Key words** Longshan Au-Sb deposit; Pyrite Re-Os dating; Zircon U-Th/He dating; Mineralization age

**摘要** 龙山金锑矿床是湘中锑-金矿集区最重要的矿床之一,因缺少适合传统放射性同位素定年的矿物,其成矿时代以往未得到很好的限定,制约了对矿床成因的认识。由于分析测试技术的进步,Re-Os 同位素定年技术得到了发展,可对热液矿床中形成的低 Re、Os 含量的硫化物进行较准确可靠的年龄测定,从而可为低温热液矿床的形成时代提供有效制约。锆石 U-Th/He 同位素定年,也是近年发展和成熟起来的定年技术,对低温热事件极其敏感,同样是约束低温成矿年龄的重要手段之一。本文采用矿床中黄铁矿 Re-Os 同位素和蚀变围岩中受成矿热事件影响的锆石 U-Th/He 同位素定年技术,对龙山金锑矿床的成矿时代进行了研究。定年结果显示:热液成因黄铁矿的 Re-Os 等时线年龄为  $195 \pm 36$  Ma,对应于印支晚期;锆石 U-Th/He 年龄为 51.2 ~ 133.3 Ma,经 Ft 校正后, U-Th/He 年龄分布于 93.78 ~ 258.29 Ma 之间,平均值为  $160.7 \pm 7.3$  Ma,对应于燕山早期。该矿床可能发生了 200 Ma 和 160 Ma 的两次成矿作用;或者矿床形成于 200 Ma 左右,但是受到了 160 Ma 左右岩浆热事件的改造,黄铁矿 Re-Os 年龄代表成矿年龄,而锆石 U-Th/He 年龄则代表第二期热事件发生的时间。无论是 200 Ma 左右一次成矿,还是另有 160 Ma 左右的成矿作用叠加,这两个年龄分别与区内两期岩浆活动的时间相当,这表明岩浆事件对驱动矿床的形成

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发挥了重要的作用。

关键词 龙山金锑矿床; 黄铁矿 Re-Os 同位素定年; 锆石 U-Th/He 同位素定年; 成矿时代

中图法分类号 P597.3; P618.51; P618.66

湘中地区是我国重要的金锑矿床产区之一,据不完全统计,已发现的金锑矿床/点超过170处(史明魁等,1993),构成华南低温成矿域的湘中锑-金矿集区(Hu and Zhou, 2012; 胡瑞忠等, 2015)。自20世纪80年代以来,许多学者针对区内的金锑矿床开展了一系列研究,但关于其成因却依然存在两种截然不同的观点:1)与岩浆岩相关的中低温热液矿床,成矿物质主要来自深部岩浆作用(刘焕品等,1985; 杨舜全, 1986; 李智明, 1993; 卢新卫, 1999; Murao *et al.*, 1999); 2)典型的层控矿床,成矿物质主要来源于沉积地层(谔锡霖等, 1983; 梁华英, 1991; 刘文均, 1992; 文国璋等, 1993; Fan *et al.*, 2004; Yang *et al.*, 2006)。争议的原因在于该类矿床的矿物组成较为简单,适合于精确开展地球化学及同位素研究的矿物相对较少,特别是缺少适合传统放射性同位素定年的矿物,极大地阻碍了对于成矿时代及成矿机制等方面的认识。锆石 U-Th/He 同位素定年是近年发展和成熟起来的一种定年技术,已在矿床学研究中取得显著进展(McInnes *et al.*, 2005a, b; Harris *et al.*, 2008; Betsi *et al.*, 2012; Li *et al.*, 2012, 2014; Liu *et al.*, 2014; Wolff *et al.*, 2015)。由于锆石 U-Th/He 同位素体系的封闭温度为160~200°C(Reiners, 2005),与低温矿床的成矿温度大致相当,低温成矿事件可以完全重置锆石 U-Th/He 同位素体系,故一般认为锆石 U-Th/He 同位素年龄可代表成矿热液冷却至锆石 U-Th/He 同位素体系封闭温度时的年龄或最后一期热液事件的年龄(Liu *et al.*, 2014)。随着分析测试水平的提高,硫化物 Re-Os 同位素体系已成为金属矿床年龄测试最为有效的技术手段之一,测试对象已从传统的辉钼矿扩展到了磁黄铁矿、黄铁矿等低 Re、Os 含量及同位素组成的矿物。近年的研究表明,Re-Os 同位素体系已能够对热液矿床中低 Re、Os 含量的硫化物矿物进行较准确的年龄测定(Stein *et al.*, 2000; Selby and Creaser, 2005; 陈懋弘等, 2007; Wang *et al.*, 2008; Liu *et al.*, 2012; 陈雷等, 2013; 邵建波等, 2014; Vernon *et al.*, 2014; Yakubchuk *et al.*, 2014; Ying *et al.*, 2014; Chen *et al.*, 2015; Huang *et al.*, 2015)。龙山金锑矿床是湘中锑-金矿集区的重要矿床之一,位于白马山-龙山成矿带的东段(图1)。前人已对该矿床的地质特征、成矿物质来源及矿床成因等开展过较多的研究(梁华英, 1989, 1991; 吴继承等, 2007; 刘鹏程等, 2008; 庞保成等, 2011),但因成矿时代尚未精确确定,导致矿床成因仍存在较大争议。本文在前人研究的基础上,通过锆石 U-Th/He 和黄铁矿 Re-Os 同位素年代学研究,确定了龙山金锑矿床的成矿时代,揭示了成矿作用与岩浆活动可能存在成因联系。

## 1 地质背景

湘中地区位于扬子地块与华夏地块的过渡部位,区域地

层具有明显的双层结构:元古界浅变质岩基底和古生界至中生界沉积盖层。其中上元古界基底浅变质岩是区内最重要的赋矿层位之一,主要由一套变质砂砾岩、板岩、变质凝灰岩组成,局部由含有基性、中酸性火山岩的碳酸盐岩和炭质板岩等组成,产出了如沃溪、渣滓溪、板溪、龙山等一些大型金锑矿床(罗献林, 1995; 梁华英, 1991; Peng and Frei, 2004; Gu *et al.*, 2012; 王永磊等, 2012)。古生界至中生界沉积盖层主要由泥质页岩、粉砂岩、灰岩、白云岩等海相、陆相及海-陆过渡相沉积岩组成,其中泥盆系细碎屑岩-碳酸盐岩系是锡矿山超大型锑矿床的主要赋矿围岩(Hu *et al.*, 1996; 彭建堂和胡瑞忠, 2001a, b; 彭建堂等, 2002; Peng *et al.*, 2003; Fan *et al.*, 2004)。该区经历了多期构造运动,形成了一系列隆起,并伴随北西向、北东向等多组不同走向断裂构造交织的构造搁架(吴继承等, 2007)。区域断裂和褶皱构造较为发育,对区内地层的产出形态及金锑矿床的分布具有明显的控制作用。尤其是印支-燕山期基底断裂的复活,切割了东西向隆起带,形成若干个次级隆起(图1a, b; 如白马山穹隆、大乘山穹隆和龙山穹隆),而这些基底断裂则起到了沟通上下构造层的作用,为含矿溶液的迁移成矿提供了通道;与此同时,印支期花岗岩体(如白马山岩体、沅山岩体、紫云山岩体等,图1a)的形成(200~230Ma; Ding *et al.*, 2006; Chen *et al.*, 2007; Wang *et al.*, 2007; Chu *et al.*, 2012; 李建华等, 2014; 刘凯等, 2014; Fu *et al.*, 2015),可能为区内金锑矿床的形成提供了热源及部分成矿物质。

龙山金锑矿床位于白马山-龙山东西向隆起带与北东向宁乡-新宁基底断裂北西向锡矿山-涟源隐伏基底断裂的交汇部位(图1b),赋存于龙山穹隆的中心部位。穹隆核部由前泥盆系基底构成,泥盆系围绕穹隆成环带状展布,石炭系-三叠系分布于白马山-龙山隆起带的南北两侧。地层整体为一套坳拉槽环境中沉积的裂谷式沉积建造(庞保成等, 2011)。断裂构造控制了龙山金锑矿床的产出位置和形态(图2)。其中,北西西向断裂多为张扭性断裂,规模较大,矿体规模也相对较大;而其他走向的断裂则多显示压性或压扭性特征,相应的矿体规模也较小。区内无大岩体出露,仅矿区西北侧发育有两条规模较小的酸性岩脉。但在龙山穹隆周边发育有大量花岗岩斑岩脉和花岗闪长斑岩脉(湖南省地质矿产局, 1988)。最新的年代学研究显示,龙山穹隆周边的酸性岩脉性主要形成于印支晚期(陈佑纬等, 2016),与湘中盆地周边大面积发育的花岗岩体形成年龄一致(Fu *et al.*, 2015),表明二者可能为同一期岩浆作用的产物。地球物理资料显示,龙山穹隆的深部可能存在大型隐伏岩体(黎盛斯, 1996; 饶家荣等, 1999)。

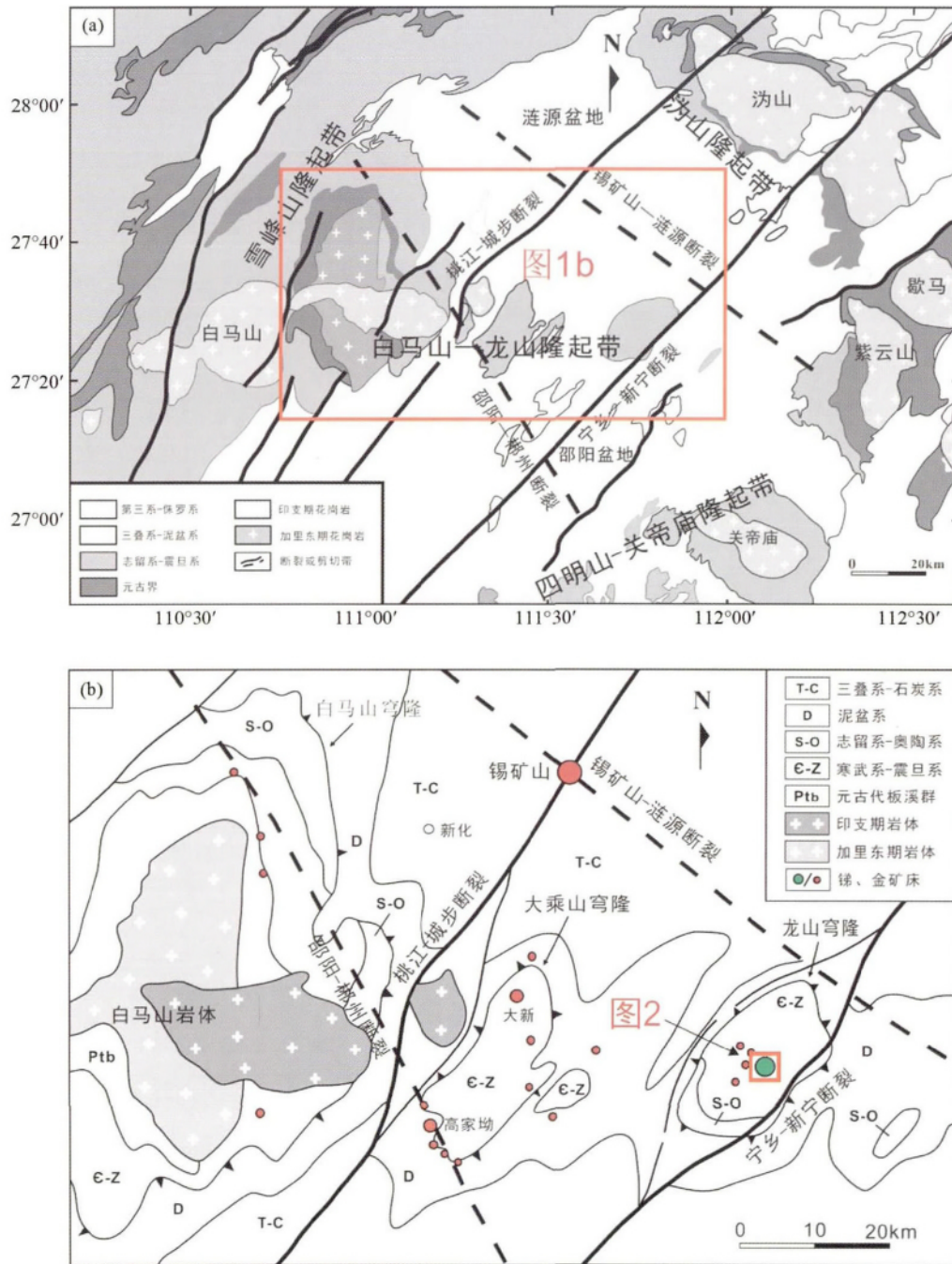


图1 湘中地区区域地质简图(a, 据湖南省地质矿产局, 1988 修改) 和白马山-龙山金锑成矿带区域地质图(b, 据李己华等, 2004 修改)

Fig. 1 Simplified geological map of the Xiangzhong district (a, after BGMRH, 1988) and regional geological map of the Baimashan-Longshan Au-Sb ore belt (b, modified after Li *et al.*, 2004)

## 2 矿床地质及样品采集

龙山金锑矿床的矿体主要赋存于龙山穹隆核部震旦系江口组的一套浅变质碎屑岩中, 该套地层自下而上分四个岩性段, 矿体主要赋存于江口组上段第一、二亚段的含砾砂质

板岩、凝灰质含砾砂质板岩中。矿体形态严格受背斜核部的断裂控制, 多呈脉状及透镜状产出(图2和图3), 具有成组分布的特点。根据矿脉展布方向, 主要分为北西向和北东向两组, 以前者为主, 主要包括1号和2号矿体, 是龙山金锑矿床早期开采的主要对象。北东向的矿体主要有5号、7号、8号和20号等矿脉, 其中5号、7号和8号矿脉是当前

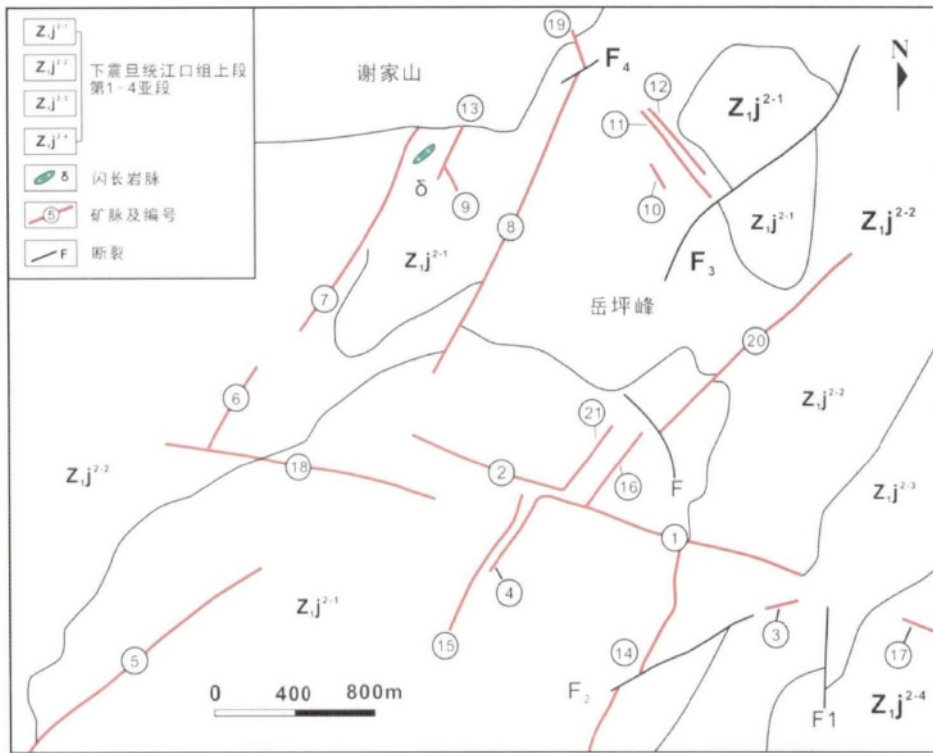


图2 龙山金锑矿床矿区地质图(据郑时干, 2006 修改)

Fig.2 Geological map of the Longshan Au-Sb deposit (modified after Zheng, 2006)

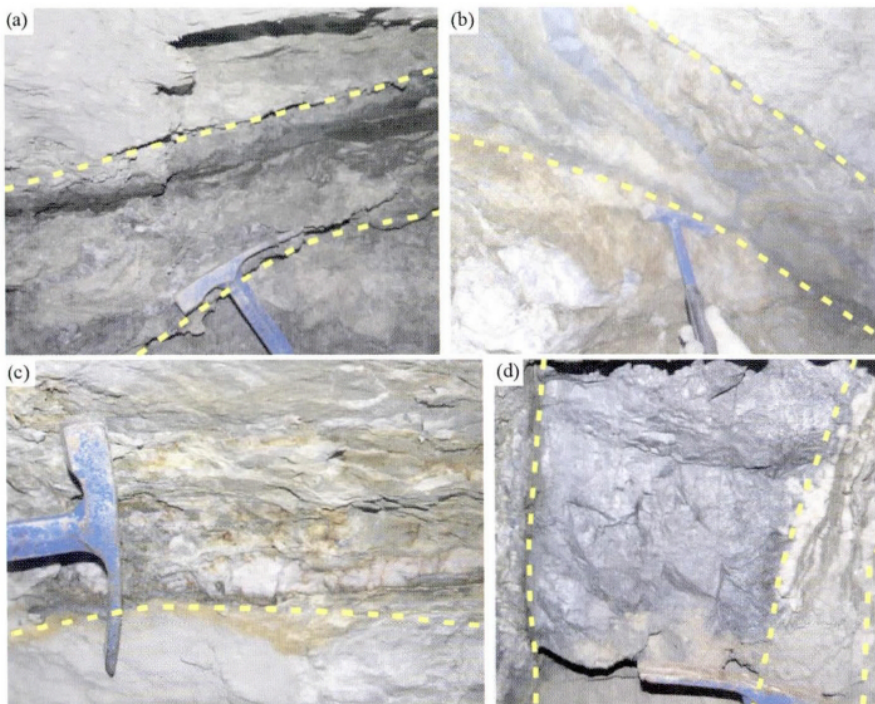


图3 龙山金锑矿床矿体形态

Fig.3 Pictures of orebodies from the Longshan Au-Sb deposit

表1 湘中龙山金锑矿黄铁矿 Re-Os 同位素分析

Table 1 Pyrite Re-Os isotopic analysis of the Longshan Au-Sb deposit in central Hunan Province

样品号	Re ( $\times 10^{-9}$ )		* Os ( $\times 10^{-9}$ )		$^{187}\text{Os}$ ( $\times 10^{-9}$ )		$^{187}\text{Re}/^{188}\text{Os}$		$^{187}\text{Os}/^{188}\text{Os}$	
	测定值	不确定度/ $1\sigma$	测定值	不确定度/ $1\sigma$	测定值	不确定度/ $1\sigma$	测定值	不确定度/ $1\sigma$	测定值	不确定度/ $1\sigma$
LS-5-1	4.4149	0.0131	0.0067	0.0001	0.0099	0.0002	3172	43	11.31	0.24
LS-5-2	0.9112	0.0068	0.0237	0.0002	0.0054	0.0000	184.9	1.9	1.739	0.005
LS-5-3	0.4410	0.0033	0.0033	0.0000	0.0017	0.0000	641.3	6.6	3.857	0.018
LS-5-4	1.4122	0.0104	0.0267	0.0002	0.0057	0.0000	254.6	2.6	1.640	0.003
LS-5-5	1.1780	0.0087	0.0306	0.0002	0.0054	0.0000	185.3	1.9	1.360	0.004

注: \* 代表普通 Os 含量

主要的开采对象。根据主要金属矿物组合特点, 矿石主要有辉锑矿-毒砂(黄铁矿)自然金、毒砂-自然金、黄铁矿-自然金三种类型(刘鹏程等, 2008; 庞保成等, 2011)。矿石矿物主要为黄铁矿、辉锑矿、自然金和毒砂, 脉石矿物主要为石英, 次为绢云母、方解石、绿泥石等。矿石结构主要为粒状结构、鳞片变晶结构, 具有脉状、网脉状、条带状、块状和角砾状构造。围岩蚀变主要为硅化、绢云母化、碳酸盐化和绿泥石化, 其中硅化和绢云母化与成矿关系密切。根据矿物组合特征, 可把成矿期划分为热液期和表生期, 其中热液期又可分为早期石英-黄铁矿(毒砂)阶段、中期自然金-黄铁矿(毒砂)-石英阶段和晚期辉锑矿-黄铁矿-石英-碳酸盐岩阶段。

本次工作的样品主要采自 5 号矿脉, 其中黄铁矿 Re-Os 同位素定年的样品为辉锑矿-黄铁矿-自然金类型的矿石, 黄铁矿成脉状或浸染状产出, 粒径多小于 1mm, 本次工作共挑选出 5 件较粗粒的黄铁矿用于 Re-Os 同位素定年; 锆石 U-Th/He 同位素定年的样品采自 5 号矿脉边部的强蚀变围岩, 采用传统的重-磁方法挑选出锆石颗粒, 然后在双目镜下挑选出自形程度高、透明好、无裂隙和包体的锆石用于 U-Th/He 同位素定年。

### 3 分析方法与结果

#### 3.1 黄铁矿 Re-Os 同位素定年及结果

黄铁矿 Re-Os 同位素测定在中国地质科学院国家地质实验测试中心 Re-Os 同位素实验室完成, 具体原理和详细分析方法见文献(杨胜洪等, 2007; 杜安道等, 2009; 李超等, 2009)。采用卡洛格管溶样, 通过蒸馏分离 Os, 用丙酮萃取并纯化 Re。样品 Re、Os 的含量及同位素组成采用 Finnigan Element 2 型高分辨电感耦合等离子质谱(HR-ICP-MS)完成。测定过程中通过监测 $^{185}\text{Re}$ 控制 Re 对 Os 测定的干扰, 通过监测 $^{190}\text{Os}$ 控制 Os 对 Re 测定的干扰。分析质量采用国家标准物质 GBW04435(HLP)和 GBW04436(JDC)进行监测, 整个实验流程空白为  $\text{Re} < 4\text{pg}$ ,  $\text{Os} < 0.7\text{pg}$ 。

5 件黄铁矿样品的 Re、Os 含量及 $^{187}\text{Re}/^{188}\text{Os}$ 和 $^{187}\text{Os}/^{188}\text{Os}$ 测定结果列于表 1。黄铁矿的 Re 含量为  $0.4410 \times 10^{-9}$

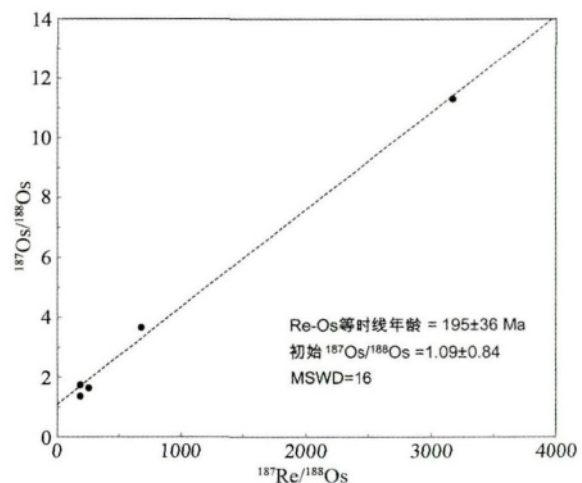


图4 龙山金锑矿床黄铁矿 Re-Os 同位素等时线图

Fig. 4 Pyrite Re-Os isotopic isochrons of the Longshan Au-Sb deposit

$\sim 4.4149 \times 10^{-9}$ , Os 含量为  $0.0067 \times 10^{-9} \sim 0.0306 \times 10^{-9}$ ,  $^{187}\text{Re}/^{188}\text{Os}$  和  $^{187}\text{Os}/^{188}\text{Os}$  比值分别为 184.9 ~ 3172 和 1.36 ~ 11.31。用 Isoplot2.06 计算获得的 Re-Os 等时线年龄为  $195 \pm 36\text{Ma}$  ( $2\sigma$ ,  $n=5$ ),  $\text{MSWD} = 16$ , 初始  $^{187}\text{Os}/^{188}\text{Os}$  比值为  $1.09 \pm 0.84$  ( $2\sigma$ ) (图 4)。本次实验结果的误差偏大, 其原因通常涉及以下几个方面: (1) 方法本身的原因, 即黄铁矿 Re-Os 同位素分析的复杂性和不确定性而所致; (2) 样品自身的原因, 如黄铁矿常发育碎裂结构, 裂隙中的充填物可导致黄铁矿样品的纯度不足而引起误差(唐永永等, 2013); (3) 黄铁矿中 Re、Os 含量偏低, 造成等时线上各点拉不开而引起误差(李超等, 2010)。由于本次测试的黄铁矿 Re 含量偏低, 其中 3 件样品的 Re 含量低于  $1 \times 10^{-9}$ , 且在等时线上分布较为集中, 这可能是导致年龄误差偏大的原因。但从 MSWD 值及拟合概率来看, 其等时线年龄是可靠的, 对于开展矿床形成时代的研究仍具有重要意义。

#### 3.2 锆石 U-Th/He 同位素定年及结果

本次工作锆石 U-Th/He 同位素定年的实验准备和测试

表2 湘中龙山金锑矿床锆石 U-Th/He 同位素分析结果

Table 2 Zircon U-Th/He isotopic dating result of the Longshan Au-Sb deposit in central Hunan Province

样品点	锆石直径* ( $\mu\text{m}$ )	$^{232}\text{Th}$ (ng)	$1\sigma$ (%)	$^{238}\text{U}$ (ng)	$1\sigma$ (%)	$^4\text{He}$ (ncc)	$1\sigma$ (%)	TAU (%)	原始年 龄(Ma)	$1\sigma$ (Ma)	$F_t$	校正后年 龄(Ma)	$1\sigma$ (Ma)
LS-5-3@1	51.75	0.375	4.5	0.587	4.1	7.795	1.5	3.9	94.2	3.7	0.5341	176.3	3.6
LS-5-3@2	49.45	0.231	4.5	0.831	4.1	14.524	6.0	7.1	133.3	9.5	0.5159	258.3	9.5
LS-5-3@3 <sup>#</sup>	42.55	0.233	4.5	0.718	4.0	5.108	28.6	28.9	54.1	15.6	0.4466	121.1	15.7
LS-5-3@4	49.45	0.101	4.5	0.198	4.0	2.042	10.9	11.5	75.3	8.7	0.5101	147.4	8.7
LS-5-3@5	55.20	0.407	4.5	0.691	4.0	4.920	8.9	9.6	51.2	4.9	0.5459	93.8	4.9
LS-5-3@6	48.30	0.209	4.5	0.976	4.0	8.110	14.6	15.1	64.7	9.8	0.5075	127.4	9.8
平均值												160.7	7.3

注: 标注<sup>#</sup>表示误差过大, 故不参与年龄计算; \* 表示将锆石的体积转化为球体后的球体直径

工作是在澳大利亚科廷大学 John de Laeter Center 同位素实验室完成的。将挑选出的锆石颗粒置于双目镜下, 选择自形程度高、透明、包体少和无裂隙的锆石颗粒, 用于 U-Th/He 同位素测定。同时记录所挑选锆石的长度、宽度等参数, 用于计算  $\alpha$  射出效应的校正系数  $F_t$  ( $F_t$  计算公式见下文; Farley *et al.*, 1996; Farley, 2002)。将每个锆石颗粒分别放入 Nb 盒内后封闭, 然后将其置入超高压真空箱内使用 1064nm Nd-YAG 激光系统在 1250°C 下提取单颗粒锆石中的 He, 并在 Pfeiffer Prisma QMS-200 质谱上进行分析。每个锆石颗粒都进行了多次去气过程直至锆石内 He 气被全部提取。采用热 Ti-Zr 吸收剂(350°C)对提出的 He 气进行纯化, 加入  $^3\text{He}$  稀释剂(纯度为 99.9%)后导入到紧邻冷 Ti-Zr 吸收剂的质谱仪内进行 He 同位素分析。 $^4\text{He}/^3\text{He}$  比值采用静态模式下的 Channeltron 探测器进行分析, HD 和  $^3\text{H}$  干扰采用质/荷比 = 1 进行校正, 采用独立  $^4\text{He}$  标准池对  $^4\text{He}$  浓度进行校正, 单个样品  $^4\text{He}$  浓度分析误差小于 1%。将去气后的锆石样品从激光室内取出转至 Parr 压溶管内, 接下来的溶解过程详见 Evans *et al.* (2005)。添加稀释剂  $^{235}\text{U}$ ,  $^{230}\text{Th}$  和 350  $\mu\text{L}$  HF 后, 在 240°C 条件下溶解 40h 左右, 采用添加同样稀释剂的 25  $\mu\text{L}$  标准溶液和未加稀释剂的 HF 空白溶液对流程空白进行监测; 然后将溶液从压力管中取出, 放置在热盘上蒸干 48h; 然后给每个样品加添 350  $\mu\text{L}$  HF, 再次放置 200°C 热盘上加热 24h 以确保氟化盐全部溶解。样品、空白和标准溶液同时在 Agilent 7700 ICP-MS 上进行  $^{238}\text{U}$  和  $^{232}\text{Th}$  含量分析。锆石 (U-Th)/He 年龄的总误差小于 5%, 采用  $F_t$  方法进行年龄校正,  $\text{Age}_{\text{corrected}} = \text{Age}_{\text{raw}}/F_t$ 。 $F_t$  校正方法由 Farley *et al.* (1996) 提出,  $F_t = 1 - 4.55\beta + 5.2\beta^2$ ,  $\beta$  为矿物颗粒表面积与体积的比值。对于四方晶系的锆石而言,  $\beta = (4L + 2W1)/(L \times W1)$ , 其中 L 为锆石颗粒长度, W1 为锆石颗粒的宽度。

6 颗锆石的 U-Th/He 同位素分析结果见表 2。其中, 锆石 LS-5-3@3 的分析结果误差较大(超过 10%), 远超出实验室分析误差, 其可靠性存疑。其余 5 颗锆石的 U-Th/He 同位素年龄为 51.2 ~ 133.3Ma。经  $F_t$  校正后, U-Th/He 年龄分布于 93.78 ~ 258.29Ma, 5 颗锆石 U-Th/He 平均年龄为  $160.7 \pm 7.3\text{Ma}$ 。已有的研究表明, 引起单个样品内 (U-Th)/He 年龄

分散的因素主要有: (1) 矿物颗粒内含有富 U 和 Th 的包体或流体包裹体 (Lippolt *et al.*, 1994); (2) 富 He 主岩的外来干扰 (Spiegel *et al.*, 2009; Danisik *et al.*, 2010); (3) 锆石颗粒的大小不同 (Reiners *et al.*, 2004); (4) 矿物颗粒内 U、Th 等不均一分布(如带状分布) (Farley *et al.*, 1996; Meesters and Dunai, 2002; Hourigan *et al.*, 2005; Danisik *et al.*, 2010; Guenther *et al.*, 2013, 2014); 与 (5) U 和 Th 衰变造成的辐射损伤 (Nasdala *et al.*, 2004; Shuster *et al.*, 2006; Shuster and Farley, 2009)。本文所获得的锆石 U-Th/He 年龄误差范围偏大, 可能是由于锆石内 U、Th 等母体同位素的不均一分布所致, 而在实验测试过程中通常难以有效地测定 U、Th 的分布状态。因此, 在处理 U-Th/He 同位素定年结果时, 常采用多颗粒年龄的平均值代表该样品的 U-Th/He 年龄。

流体包裹体研究表明, 龙山金锑矿床的形成温度位于 165 ~ 230°C 之间 (梁华英, 1991; 马东升等, 2003), 高于锆石 U-Th/He 同位素体系的封闭温度 ( $180 \pm 20^\circ\text{C}$ ; Reiners, 2005), 成矿热事件可完全重置锆石的 U-Th/He 同位素体系进而记录成矿事件发生的时间。目前 U-Th/He 同位素定年已在金矿床、斑岩型铜-金矿床等研究中取得了较好的应用 (McInnes *et al.*, 2005a, b; Harris *et al.*, 2008; Betsi *et al.*, 2012; Li *et al.*, 2012, 2014; Cabral *et al.*, 2013; Zeng *et al.*, 2013; Liu *et al.*, 2014), 可对成矿时代、矿化持续时间及成矿后的抬升剥蚀和保存情况进行较好的制约 (陈文等, 2010)。本次工作测试的锆石选自矿脉边部强蚀变的围岩, 其内的锆石 U-Th/He 同位素体系受到成矿事件的影响而发生重置, 故其 U-Th/He 同位素年龄可为研究龙山金锑矿床的形成时代提供约束。

## 4 讨论

由于缺少适合用传统放射性同位素法定年的矿物, 龙山金锑矿床的形成时代一直未得到有效制约。梁华英 (1991) 根据构造与成矿关系及龙山地区出露的燕山期酸性岩脉, 推测龙山金锑矿床的形成时代可能为燕山期, 同时认

为还存在加里东期成矿的可能。史明魁等(1993)通过石英流体包裹体 Rb-Sr 等时线法获得了  $175 \pm 27$  Ma 的年龄,并认为龙山金锑矿床的形成与区内燕山早期的岩浆活动有关。然而最新的研究结果显示,龙山地区的酸性岩脉主要形成于印支晚期(217~220 Ma; 陈佑纬等,2016),并非前人推测的燕山期。事实上,印支晚期的酸性岩体/脉在湘中地区普遍发育,且与金锑矿床呈现出密切的空间关系,近年来在这些酸性岩体/脉周边发现了一系列金锑矿床/点,如产于白马山岩体边部的古台山金矿床(戴长华,2000a)、青京寨金矿床(余建国,1998;戴长华,2000b)、铲子坪和大坪金矿床(李华芹等,2008)、沩山岩体边部的太平金矿床(卢新卫,1999)以及紫云山岩体边部的铃山和马鞍山金矿床(戚学祥,1998;王滨清,2005)等。刘继顺(1996)的研究表明,湘中地区至少有 27 个金锑矿床/矿点,在其 100 米范围内见有酸性岩脉或小岩珠分布,一些地方这些岩脉或小岩珠甚至为矿体的直接围岩,这表明湘中地区酸性岩体/脉与金锑矿化之间存在密切的空间联系。前人通过对赋矿岩脉年龄的测定来限定金锑矿床的形成时代,如通过对廖家坪金锑矿床、符竹溪锑金矿床和板溪锑矿床赋矿酸性岩脉年龄(钾长石 K-Ar)的测定,获得相应矿床的年龄分别为 200 Ma、209 Ma 和 194~204 Ma(肖启明等,1992;姚振凯和朱蓉斌,1993;彭建堂和胡瑞忠,2001a)。李华芹等(2008)测定的位于白马山岩体边部的铲子坪和大坪金矿床的石英流体包裹体 Rb-Sr 等时线年龄分别为  $205.6 \pm 9.4$  Ma 和  $204.8 \pm 6.3$  Ma。上述研究表明,湘中地区在印支晚期(~200 Ma)可能存在一期与同期岩浆活动密切相关的金锑成矿作用。

本次工作获得黄铁矿 Re-Os 同位素等时线年龄为  $195 \pm 36$  Ma,与龙山地区出露的酸性岩脉及湘中盆地周边的印支期花岗岩体的形成时代相近。但蚀变围岩内受成矿热事件影响的锆石的 U-Th/He 年龄为  $160.7 \pm 7.3$  Ma,远小于区内岩浆活动的时间,也与黄铁矿 Re-Os 同位素年龄相差较大。由于锆石 U-Th/He 同位素体系的封闭温度较低,易受到后期热事件的改造或重置。因此,锆石 U-Th/He 年龄所代表的地质意义通常分为两种情况:单一冷却过程(Single cooling)对应的锆石 U-Th/He 年龄代表了热事件冷却至 U-Th/He 同位素体系封闭温度时的时间,即热事件晚期的年龄;复杂冷却过程(Multiple cooling)锆石 U-Th/He 同位素体系可能经历了多次重置过程,对应的 U-Th/He 年龄则代表了 U-Th/He 同位素体系最后一次重置发生的时间,即最后一次热事件发生的时间(Liu *et al.*, 2013)。那么,我们假设锆石 U-Pb 同位素体系封闭温度为  $900^\circ\text{C}$ (Lee *et al.*, 1997; Cherniak and Watson, 2001),区内岩浆活动的最小年龄为 210 Ma;锆石 U-Th/He 体系的封闭温度为  $200^\circ\text{C}$ , U-Th/He 年龄为 160 Ma,那么成矿热液的冷却速率约  $14^\circ\text{C}/\text{Myr}$ ,远小于典型岩浆热液矿床的冷却速率(McInnes *et al.*, 2005a; Harris *et al.*, 2008; Li *et al.*, 2012, 2014)。这种不匹配性的原因可能为:(1)区内印支晚期岩浆活动后还存在一期热事件,这一热事件完全重

置了锆石 U-Th/He 同位素体系;(2)成矿事件发生于深部,围岩与成矿热液达到了热平衡,且高于锆石 U-Th/He 同位素体系的封闭温度。已有的研究结果更支持第(1)种可能,如前人在研究湘中地区的印支期花岗岩体时发现,在白马山、沩山、紫云山等印支期岩体内存在少量燕山早期(170~180 Ma)的锆石(Ding *et al.*, 2006; Chen *et al.*, 2007; 刘建清等,2013),表明湘中地区可能存在燕山早期的岩浆活动,但地表出露不显著。另一方面,许多研究表明湘中地区也存在燕山早期的金锑成矿年龄,如形成于 155 Ma 左右的锡矿山锑矿床(Peng *et al.*, 2003)以及龙山地区最近发现的谢家山锑金(钨)矿床也形成于 160 Ma 左右(谢桂青等,未发表数据)。上述研究表明,燕山早期的岩浆活动促使湘中地区发生了又一次金锑成矿事件。

因此,我们认为龙山金锑矿床极可能:(1)发生了 200 Ma 和 160 Ma 的两次成矿作用,矿床是这两期成矿作用叠加的产物;或者(2)矿床形成于 200 Ma 左右,但是受到了 160 Ma 左右岩浆热事件的改造,黄铁矿 Re-Os 同位素年龄代表成矿事件的年龄,而锆石 U-Th/He 同位素年龄则代表了第二期热事件发生的时间。

龙山金锑矿床的成因一直备受争议,争议的焦点在于成矿作用是否与岩浆活动有成因联系(梁华英,1991;马东升等,2002,2003;刘鹏程等,2008;庞保成等,2011)。我们的研究表明,无论该矿床是 200 Ma 左右一次成矿,还是另有 160 Ma 左右的成矿作用叠加,这两个年龄都分别与区内两期岩浆活动的时间相当,表明金锑矿床与岩浆活动之间存在明显的时间相关性,这表明岩浆事件对驱动矿床的形成发挥了重要的作用。

## 5 结论

(1) 龙山金锑矿床可能发生了 200 Ma 和 160 Ma 的两次成矿作用,矿床是这两期成矿作用叠加的产物;或者矿床形成于 200 Ma 左右,但是受到了 160 Ma 左右岩浆热事件的改造,黄铁矿 Re-Os 同位素年龄代表成矿事件的年龄,而锆石 U-Th/He 年龄则代表了第二期热事件发生的时间。

(2) 无论该矿床是 200 Ma 左右一次成矿,还是另有 160 Ma 左右的成矿作用叠加,这两个年龄都分别与区内两期岩浆活动的时间相当,这表明岩浆事件对驱动矿床的形成发挥了重要的作用。

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