

地表自然过程排汞研究进展及展望*

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摘要 地表自然过程排汞包括了自然源排汞过程和先前排放的汞沉降到地表后的再排放过程。已有证据显示, 地表自然过程向大气的排汞量可能远大于人为活动直接向大气的排汞量。准确确定自然过程汞的释放通量, 不仅对正确评价目前减少人为活动向大气排汞对全球环境汞污染的影响程度具有重要的意义, 而且可为全球大气汞的减排政策的制定提供重要科学理论依据。本综述通过对国内外地表自然排放源相关文献的调研分析发现: 由于缺少可靠的观测技术、对地表与大气间汞交换过程和机理的准确认识及大气汞沉降对地表自然排汞过程影响的认识还不清楚, 因此目前还难以准确估算地表自然过程向大气的排汞量。近年来, 随着技术手段的进步, 已具备了开展地表自然排汞及先前排汞沉降后的再释放过程、机理和通量研究的可行性。地表自然过程排汞的研究将是汞的生物地球化学循环演化研究领域的前沿之一。

关键词 汞; 自然来源; 交换通量; 微气象方法; 汞同位素示踪

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Abstract Earth surface natural mercury (Hg) emission includes the Hg emission from natural sources and the re-emission from previously deposited Hg. It was demonstrated that the total amount of natural Hg emission from earth surface could be far larger than that of the direct emission from human activities. It is of great importance to quantify the natural Hg emission, not only for the evaluation of the effects of reducing anthropogenic Hg emission on the global environment, but also for providing guidance on the policy-making in global Hg reduction strategy. Due to the lack of reliable methodologies in quantifying the Hg emission from natural sources, we still don't have a clear understanding about the processes and mechanisms of Hg exchange between earth surface and atmosphere as well as the effects of atmospheric Hg deposition on the processes of earth surface natural Hg emission and thereby we are hard to accurately quantify the Hg emission from natural sources. However, with the development of new technologies, we have the feasibility to study the processes, mechanisms, and fluxes of Hg emission from natural sources. To quantify the Hg emission from natural sources would be one of the research frontiers about the biogeochemical cycling of Hg in the environment.

Key words mercury; natural source; exchange flux; micrometeorological method; mercury isotope tracer

由于甲基汞极强的毒性, 汞被各国政府和国际

卫生组织定义为环境中最有害的重金属元素之一。

20世纪90年代, 在北美和北欧偏远地区清洁湖泊中鱼体汞含量已超出世界卫生组织建议的食用标准

($0.5 \text{ mg} \cdot \text{kg}^{-1}$)¹, 发现人类活动和自然过程向大气

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排放的汞经长距离迁移后在这些地区的沉降是引起这类汞污染的主要原因 (Lindqvist et al., 1991)。由此,科学家们认识到排放到大气中的气态汞可在区域和全球范围内随着大气环流进行长距离的传输,并沉降到远离汞排放源的地区,导致生物体内汞和甲基汞含量的增加。因此,汞被认为是一种通过大气进行长距离跨国界传输的全球性污染物 (Lindqvist et al., 1991)。

由于特殊的物理化学性质,汞是在大气中唯一以气态形式存在的重金属元素。大气中汞主要分为三类,即气态原子汞或单质汞 (Hg^g)、活性气态汞和颗粒态汞 (Schroeder & Munthe, 1998)。 Hg^g 是大气汞的主要赋存形态,占大气总汞的 95%以上。由于 Hg^g 在水中的溶解度很低,且与大气中其他组分(如强氧化剂 O_3 、 H_2O_2 、 OH 和 NQ 自由基)的化学反应速率相对缓慢,因此 Hg^g 在大气中的居留时间很长(0.5~2年),能迁移数万公里 (Lindberg et al., 2007)。活性气态汞化合物在水中的溶解度比 Hg^g 要高,它们很易通过湿沉降和干沉降过程被带到表生生态系统中,所以在大气中的居留时间最短(数小时到几天),只能被迁移数十到几百公里。颗粒态汞在大气中的居留时间取决于颗粒物的粒径,一般在数天到几个月,在大气中的迁移距离为几百至上千公里。不同形态汞在大气中的迁移范围不同,对生态环境的影响范围也有显著差异,只有气态单质汞才是全球性污染物。

联合国环境规划署(UNEP)已正式将汞纳入环境外交。在 2009 年 2 月召开的 UNEP 第 25 届理事会会议上,与会各国环境部长们同意就起草一项旨在控制全球汞排放的国际公约展开谈判,这一公约有望在 2013 年正式启动。我国正面临着削减汞排放的巨大国际压力。

研究表明,汞排放主要有 2 种途径,一是人为活动向大气排放汞,二是自然过程向大气排汞 (Lindberg et al., 2007)。人为活动向大气排放汞主要包括燃煤、垃圾焚烧、氯碱生产、水泥生产、金属冶炼与加工等 (Wu et al., 2006)。自然过程排汞包括了自然源排汞过程和先前排放的汞沉降到地表后的再排放过程 (Lindberg et al., 2007)。主要的自然源排汞过程有:土壤、水体、植被、火山活动、森林火灾和地壳去气作用向大气的排汞 (Schroeder & Munthe, 1998)。先前排放的汞沉降到地表后的再排放过程主要指自然和人为活动向大气排放的汞在大气中被

转化成二价汞,通过大气干湿沉降返回到地表,经各种物理、化学和生物作用还原成气态原子汞再释放到大气的过程,由于目前无法区分自然排汞与先前沉降汞的再释放,因此将其统称为自然排汞 (Lindberg et al., 2007)。

目前全球关于削减汞排放重点放在加强人为活动对汞污染防治问题研究。近年来我国科学家加强了人为活动对汞污染防治问题研究 (Wang et al., 2010; Yan et al., 2009),政府部门出台了更多环保措施,在全球汞污染治理中也承担了更多责任。环境保护部门加强了对汞的生产、使用、进口、出口及加工的管理,汞矿开发受到严格限制。同时,我国大力发展洁净煤技术和新能源产业,由煤炭燃烧造成的汞污染也有所下降。然而,随着近年来关于全球汞污染问题研究的不断深入,研究人员开始意识到在加强控制人为活动向大气排汞的同时,应高度关注自然过程即自然源排汞过程和先前排放的汞沉降到地表后的再排放过程对大气排汞的贡献。

1 自然过程与人为活动向大气排汞量的估算与比较

研究显示,尽管全球范围内人为活动造成的大气汞排放在逐年降低,但大气中汞含量并没有降低。例如,自 1995 年,科学家们开始在一些全球背景观测点如加拿大北极的 Aler 观测站(全球大气本底观测站)(Steffen et al., 2005)、爱尔兰的 Mace Head 观测站和德国的 Zingst 站等 (Kock et al., 2005) 展开大气汞含量的长期连续观测,他们的观测结果一致地发现(图 1):从 1995—2004 年大气汞的背景含量保持稳定。加拿大 Aler 站自 1995 年至今,大气汞含量的年平均值一直稳定在 $1.7 \text{ ng} \cdot \text{m}^{-3}$ 。但期间,全球人为活动向大气的排汞量却有较大的变化。

如表 1 所示,1995—2000 年,全球人为活动年排汞量为 2180,但到 2005 年减少到 1927,排汞量减少了 11.6%。当然,各区域汞排放量变化趋势不同,如非洲、大洋洲和北美洲排汞量在降低;亚洲、欧洲和南美洲排汞量在增长。然而全球人为活动向大气排汞量在此期间的显著降低在全球背景站的长期观测结果中却没有显现出来。这可能说明,自然过程(包括先前排汞沉降后的再释放)向大气的排汞量可能远大于人为活动直接向大气的排汞量。再如,美国科学家通过在中国大陆的下风向对大气汞的观测及结合模型的研究表明,2004 年从中国大陆

3 地表自然过程排汞通量的估算研究进展及展望

对自然界地表向大气的汞释放通量的估算基于对选择区域进行测定基础上,构建通量结果与各种影响通量变化因子的关系式,推广到大尺度的应用计算。在这一过程中,区域的选择、测定方法的选择尤为重要。区域的选择可以按照相似的地质背景、地表类型等分类方法进行,以减少误差。

3.1 地表汞排放通量观测方法的比较与评价

实际上,自20世纪80年代末,科学们就开始致力于地表向大气排汞通量观测(Schroeder et al., 1989; Xiao et al., 1991)。目前,国际上测定地表界面与大气间的汞交换通量方法包括通量箱法和微气象学法,表3列出了各种方法的适用范围和优缺点。由于成本极低,绝大多数的研究组都是利用动力学通量箱法(dynamic flux chamber, DFC)测定地表界面与大气间汞的交换通量,如美国Gustin课题组(Erickson et al., 2006)、Lindberg课题组(Lindberg et al., 2005)和Capri课题组(Capri & Chen, 2002);加拿大的Poissan课题组(Poissant et al., 2004)、Schroeder课题组(Schroeder et al., 2005)和St. Louis课题组(Graydon et al., 2006);欧洲的Ebinghaus课题组(Bahmann et al., 2006)、Ferrara课题组(Ferrara et al., 1998)、Horvat课题组(Kooman & Horvat, 2010)和Lindqvist课题组(Xiao et al., 1991)、韩国的Kim课题组(Kim et al., 2001)以及国内笔者课题组(Feng et al., 2005; Fu et al., 2008)和王定勇课题组(Wang et al., 2006)。该方法封闭一小块地表,在测定封闭地表上空大气汞含量随时间变化速率的基础上,计算地表与大气间的交换通量(Xiao et al.,

1991)。而各个研究组使用的通量箱,从材质、外形到尺寸都有不同的设计(表4),外形有长方形、半球形、半圆柱形等,使用条件如抽气流速也从 $1.5\sim15\text{ L}\cdot\text{m}^{-2}\cdot\text{min}^{-1}$,材质有石英玻璃、聚碳酸酯、不锈钢、聚四氟乙烯(teflon)、聚丙烯酸、树脂玻璃等,测定区域面积为 $0.03\sim0.16\text{ m}^2$,通量箱体积为 $1\sim32\text{ L}$ 差别较大。

图2展示了目前最常用的2种通量箱的设计图,左通量箱的气流方向为由右到左,右通量箱从四周的进气孔进气,由上面箱盖中心出气。图2b-d是根据计算流体动力学模拟得到的通量箱内的气流场、气流流速向量分布和箱内空气汞含量分布。从图2b可以看出,两种通量箱内的气流场均为不稳定变速流,即在每个截面上风场都不相同,这样使得气流流速在整个箱测定区域内不是均匀分布(图2c),引起通量箱中汞含量的空间分布也截然不同(图2d),最终造成了在通量箱测定区域的不同位置有不同的通量。

在影响地表与大气汞交换通量的因素中,气流在地表的摩擦速率是控制汞交换过程的一个基本驱动力。对于传统的通量箱,进入通量箱的流速不同会导致测得的通量结果出现巨大差异,且不能建立流速变化与测得通量结果间的关系,不能通过流速校正对通量结果进行统一校正。因此,传统通量箱设计的不同造成每个课题组所报道的研究结果不具有可比性(Eckley et al., 2010),也不能综合利用各课题组的研究结果实现对大面积地表与大气汞交换通量的精确估算。

采样方法的缺限,是制约准确定量估算地表向大气排汞通量的最主要原因。通过上述阐述,我们

表3 通量箱法和微气象学方法测定界面大气汞通量的比较

Table 3 Comparisons between dynamic flux chamber method and micrometeorological methods for measuring Hg exchange fluxes between surface and air

方法	概念和理解	适用系统	优点	缺点
箱式法	根据密闭箱内大气汞含量的变化确定界面大气间的汞交换通量	农田、草地等低矮植被生态系统	方便、简单、易操作、价格低廉	空间代表性差、不能长时间测量、受人为小环境影响、难以真实反应界面汞交换
微气象学 涡旋相关法 (eddy covariance method, EC)	根据大气湍流状况和大气汞浓度推导出界面大气间的汞交换通量	适用范围较广	对地表无干扰;长期连续大面积监测	要求信号响应足够快速的传感器、检出限非常低的检测器,成本高、难维护,目前无法实现
微气象学 弛豫涡旋积累法 (relaxed eddy accumulation, REA)	根据大气湍流状况和大气汞浓度推导出界面大气间的汞交换通量	适用范围较广	对地表无干扰;长期连续大面积监测;对仪器时间分辨率和检测限要求较低,易操作,结果可靠	采样设备较复杂
微气象学 通量梯度法 (flux gradient methods)	根据大气湍流状况和大气汞浓度推导出界面大气间的汞交换通量	适用范围较广	对地表无干扰;长期连续大面积监测;易操作	很难精确测定非常小的含量梯度

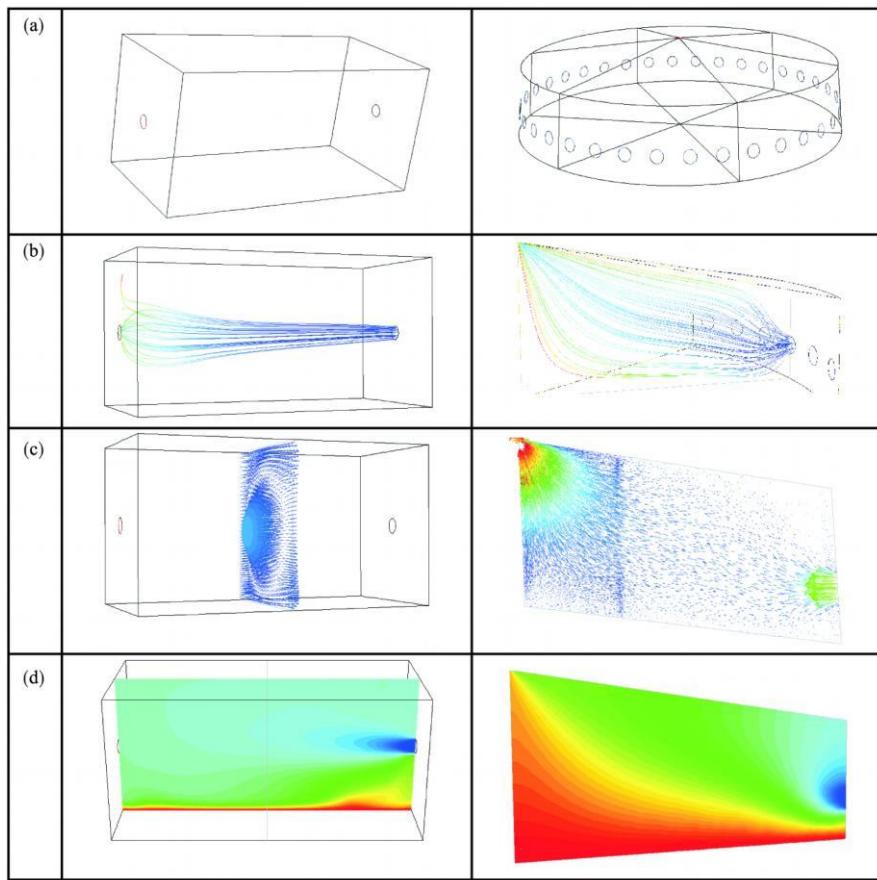


图 2 2种常用设计方案通量箱内气场及汞含量分布情况(均为任意单位)

Fig. 2 Distribution of flow rate and Hg concentrations in two commonly used flux chambers

(a)通量箱的尺寸; (b)通量箱内气流场; (c)气流流速矢量分布; (d)通量箱内空气汞含量空间分布; 通量箱底部为土壤(据 Eckley et al., 2010)。第1种设计, 通量箱的进气口和出气口只有1个, 分别在通量箱的2个侧面; 第2种通量箱设计, 进气口36个, 均匀分布在圆柱体的面上, 出气口只有1个, 分布在通量箱顶部。

表 4 各课题组采样通量箱的设计及使用材料情况

Table 4 Designs and materials of flux chambers used by different research groups

测量面积 (m ²)	通量箱 体积 (L)	采样流速 (L·m ⁻² ·min ⁻¹)	通量箱材料	土壤汞含量 (μg·g ⁻¹)	汞通量 (ng·m ⁻² ·h ⁻¹)	参考文献
0.073	11.2	1.5(7.5)	聚碳酸酯	0.072	0.4±0.5	Kuiken et al., 2008
0.13	10	1.5(6.7)	不锈钢/聚四氟乙烯	~0.03	3.0±2.2	Poissant & Casimir 1998
0.067	9.3	1.5(6.2)	聚碳酸酯	0.025~0.047	5.69±5.79	Gabriel & Williamson, 2008
0.03	2.2	1.5(1.5)	聚四氟乙烯	5	~0±45**	Zhang et al., 2002
0.036	1.5	1.5(1)	聚碳酸酯	0.013*	2.1	Eriksen & Gustafsson 2006
0.13	25	1.8(13.9)	聚四氟乙烯	0.6~30.9	200±90	He et al., 1998
0.16	32	2.0(16)	不锈钢	NA	1.4±0.5*	Schroeder et al., 1989
0.063	3	2.31(1.3)	聚碳酸酯	0.061*	0.7~35	Eckley & Branfireun, 2008
0.12	24	3.2(7.5)	丙烯酸	0.16±0.01*	0.4±0.4	Magarelli & Foster 2005
0.12	24	5(4.8)	聚四氟乙烯	0.47±0.08*	2.7±0.5	Capri & Lindberg 1998
0.12	24	6(4)	塑胶玻璃	NA(水)	0.1~44.0	Fernara & Mazzola 1998
0.12	24	6(4)	聚四氟乙烯	<0.1	1.4±1.4	Zhang et al., 2001
0.031	1	6.5(0.15)	聚碳酸酯	<0.1	2.0±4.1	Nacht & Gustafsson 2004
NA	2	6.7(0.3)	聚碳酸酯	0.02~3	<1~70	Gustafsson & Stamenkovic 2005
0.031	1	10(0.1)	聚碳酸酯	0.02~9060	0~27600***	Engle et al., 2001
0.06	4.7	15(0.31)	石英	614*	862±807	Wang et al., 2005

据 Eckley et al., 2010。()内数据为空气滞留时间 (min); 梅通量值为 mean±σ。* 数据来源于文章中采样的一个地点或者一种材料; ** 黑暗条件下采样数据; *** 数据来源于多个采样地点。NA为无数据。

已经完全理解现有通量箱法无法准确定量地表与大气汞交换通量的根本原因,对此,可以通过改进设计,逐步改善动力学通量箱法,以获得地表与大气间汞交换通量的定量结果。因此,建立新型动力学通量箱,定量测定地表与大气间汞的交换通量是今后研究工作的重点之一。

3.2 对地表与大气间汞交换过程和机理的研究现状

影响陆地生态系统(包括裸露土壤、沙漠和戈壁、草地、农田、森林、淡水湖泊和河流)向大气释放汞的因素非常多。总体来说,主要受环境介质的物理化学性质(土壤和水体汞含量和形态、土壤湿度和空隙度、土壤有机质、水体可溶性有机质、土壤和水体温度等)、环境和气象因素(如地形、植被覆盖、光照强度、风速、气温、大气氧化物)等多种因素的共同影响,表现出极强的时间和区域性变化特征。如前所述,近年来各国学者越来越关注陆地生态系统汞释放通量的研究。总体来看,这些研究主要集中于裸露土壤、湖泊河流水体、草地、湿地和森林等方面。地表和大气汞交换通量表现为一个双向的动态过程,即地表可作为大气汞的源,也可以是大气汞汇。地表汞释放的强度具有明显的区域分布特征,主要表现为:①汞矿化带土壤和岩石的释放通量显著高于相对清洁的背景地区,而这很可能与地表基质的汞含量有关(Gustin et al., 2000; Coolbaugh et al., 2002; Schröder et al., 2005; Wang et al., 2005)。②人为耕作活动能导致汞释放量的增加,如贡嘎山地区农田汞释放通量最高达到了 $132 \text{ ng} \cdot \text{m}^{-2} \cdot \text{h}^{-1}$,而同期裸露土壤的汞释放通量则为 $20 \text{ ng} \cdot \text{m}^{-2} \cdot \text{h}^{-1}$ (Fu et al., 2008)。③裸露地表的汞释放通量明显高于有植被覆盖的森林地区地表汞释放通量,植被冠层能降低透射到地表的光照强度,而光照强度是影响土壤汞释放通量的一个关键因素(Carpi & Lindberg 1997; Feng et al., 2005; Moore & Carpi 2005)。研究指出,干湿沉降到地表的氧化态汞能被还原并重新释放大气中(Erickson et al., 2005; Xin et al., 2007^a)。

3.2.1 地表裸露土壤向大气排汞的机理和影响因素已较清楚 地表裸露土壤汞释放通量受多种因素相互作用的影响,包括:

(1) 地表基质的汞含量。研究表明,地表汞释放通量与基质汞含量呈对数正线性关系(Coolbaugh et al., 2002; Schröder et al., 2005),这说明随着地

表基质汞含量的升高,汞释放通量则会出现明显升高。

(2) 辐射强度。大量的野外观测发现,地表汞释放通量和辐射强度之间存在显著的正线性相关关系(Carpi & Lindberg 1997; Poissant & Casimir 1998; Boudah et al., 2000; Zhang et al., 2001; Gustafsson et al., 2002^{a, b}; Feng et al., 2005; Fu et al., 2008)。而实验室的研究表明(Bahmann et al., 2004; Moore & Carpi 2005),在恒定温度和其他环境条件下,辐射强度的增加确实能显著提高土壤汞的释放通量。研究还表明(Moore & Carpi 2005),在全部辐射的波谱内,只有紫外波段的辐射才是最主要的影响因素,而其他波段辐射的影响很小。这表明,土壤中二价汞的光致还原作用是土壤向大气排汞的驱动力(Feng et al., 2005)。

(3) 土壤和大气温度。Moore 和 Carpi(2005)通过室内试验研究指出,土壤温度的升高,能够促进汞释放通量增强。此外,Fu等(2008)通过对避光条件下汞释放通量日变化的研究指出,土壤温度可能也是一个重要影响因素。在特殊的环境条件下,比如森林(辐射强度波动大且地表微生物活动较强),温度对汞释放通量的影响就比辐射强度明显(Choi & Holsten, 2009)。

(4) 风速。风速对地表汞释放通量的影响主要是通过改变地表湍流状况实现的。许多野外观测发现,土壤汞释放通量与风速呈正相关关系,说明风速的增强能够促进土壤汞的解吸附过程,从而促进汞的释放(Lindberg et al., 1999; Wallischger et al., 1999; Gillis & Miller, 2000)。不过,风速的促进作用可能受到其他条件的制约。研究发现,高汞土壤汞释放通量对风速的响应较为明显,而风速对低汞含量土壤的汞释放通量影响就很小(Gustin et al., 1997; Poissant et al., 1999)。不过,目前所采用的动力学通量箱法对研究风速的影响作用是有很大局限的。有研究表明,通量箱法测定结果和实际值的偏差会随风速的增大而增大(Gillis & Miller, 2000)。

(5) 大气汞含量。研究指出,土壤汞释放通量取决于壤中气和大气汞含量的梯度(Zhang et al., 2002),这意味着土壤的汞释放通量会因大气汞含量的升高而受到抑制(Gillis & Miller, 2000)。Xin 和 Gustin(2007)的研究发现低汞含量的土壤会随大气汞含量的升高而出现线性降低,甚至出现明显沉降。不过,随着人为活动汞释放的增加,大气汞含量

会出现显著的升高和波动, 因此会导致个别沉降事件的发生。比如, Wang 等 (2007) 和 Fu 等 (2008) 对务川汞矿区和贡嘎山地区的研究表明, 人为汞释放强度的增强, 会导致明显的大气汞向地表的沉降。

除上述因素外, 土壤汞释放通量还受到其他多种因素的影响。比如, 土壤水分和 pH 的升高都能一定程度上促进土壤汞的释放 (Lindberg et al., 1999; Gillis & Miller, 2000; Johnson et al., 2003; Gustafsson & Stamenkovic, 2005), 而土壤有机质则对汞释放通量具有一定抑制作用 (Maucclair et al., 2008)。此外, 大气氧化还原特性 (如 O₃ 浓度)、土壤铁锰离子和铝氧化物含量、土壤类型对汞释放通量也有一定影响。总的来说, 目前对裸露土壤向大气排汞机理和影响因素的认识已较为清楚。

3.2.2 目前对植被与大气间汞的交换过程和机理认识还不清楚 Lindberg 等 (1998, 2002) 采用微气象梯度法测定森林汞释放通量范围为 8~37 ng·m⁻²·h⁻¹, 表现为大气汞的净源; 同时森林汞释放是由植物释放占主导, 而土壤排汞的贡献很小。在此基础上, Shetty 等 (2008) 采用土壤溶液的汞含量和植被的蒸腾速率估算了亚洲地区植被每年向大气的排汞通量, 结果显示, 亚洲的植被汞释放通量是大气汞的一个重要来源, 年均释放量约为 630 t·a⁻¹。Quar 等 (2008) 采用生物源排放量推估模型估算的我国植被每年的大气排汞量约为 79~177 t·a⁻¹。由于对植被与大气间汞的交换机理还认识不清, 上述研究结果受到一些置疑。例如, 利用通量袋法测定结果表明, 植物叶片是大气汞的汇, 平均沉降通量为 0.4~1 ng·m⁻²·h⁻¹ (面积为叶片面积) (Zhang et al., 2005; Poissant et al., 2008)。另有研究指出, 植物叶片对大气汞的吸收可能是植物叶片中汞的唯一来源 (Erickson & Gustafsson, 2004; Greger et al., 2005; Bushay et al., 2008; Stamenkovic & Gustafsson, 2009), 这意味着植被可能是大气汞的汇。Hartman 等 (2009) 利用模型估算了美国大气汞向叶片的沉降通量为 15 t·a⁻¹。

由于目前对植被与大气间汞交换机理的认识还很欠缺, 无法判断植被是大气汞的源还是汇。全球大部分的陆地表面为植被所覆盖, 显然只有深入了解植被与大气间汞的交换机理和交换通量, 才能正确估算全球地表向大气的自然排汞通量。因此, 汞在土壤、植物、大气间的迁移转化规律将是今后研究工作的重点之一。

3.2.3 大气汞沉降对地表自然排汞过程的影响还不清楚 地表的汞除了来源于基岩风化过程 (地质源) 外, 大气沉降源源不断地向地表输入汞 (Nater & Grigal, 1992)。有研究表明, 大气沉降的汞主要在土壤有机质层即土壤 A 层蓄积 (Fitzgerald et al., 1998; St. Louis et al., 2001)。在远离自然汞富集地区如汞矿化带和人为汞污染源区, 土壤中汞大部分来源于大气汞的沉降 (Werner et al., 2006)。由于工业化过程人为活动向大气排放大量的汞, 目前全球土壤中汞含量与工业革命前比有显著的增长 (Mason et al., 1994)。但是, 除了在地表土壤中蓄积外, 有多少汞重新释放到大气无法确定。Erickson 等 (2005) 在美国沙漠地区开展了大气沉降汞有多少会重新释放到大气的实验研究, 在 4 m² 沙漠土壤中加入一定量的只含单一汞同位素 ¹⁹⁸HgCl₂ 溶液, 利用动力学通量箱法测定土壤向大气的排汞通量, 尤其是测定了过剩 ¹⁹⁸Hg (样品中 ¹⁹⁸Hg 减去自然界中 ¹⁹⁸Hg 的量) 的释放通量, 结果表明, 只有 6% 的加入汞在 1 年时间内重新释放到大气中。由于沙漠土壤中有机质含量很低, 加入的汞会随溶液渗透到光照无法达到的土壤深部, 因此加入的汞无法被光还原为单质汞, 从而无法释放到大气中。这一研究结果可能在富含有机质的土壤并不适用, 因为土壤表层有机质会吸附大气降水中的汞, 使其在表层土壤富集。表层土壤的汞容易被光还原, 重新释放到大气。但目前有关大气汞沉降对其他类型地表排汞的影响还不清楚。弄清大气汞沉降对不同地表类型向大气排汞通量的贡献量, 将有助于认识新沉降汞的再释放占自然排汞通量的比例, 从而为人为活动减排政策的制定提供重要的理论依据。

4 测定地表向大气自然排汞通量技术手段的研究进展与展望

微气象方法是近年来发展起来的测定地表界面与大气间微量气体交换通量的新方法。微气象方法 (micrometeorological methods) 是通过测定界面附近的湍流状况和大气微量气体含量的垂直变化进而推导出界面大气间的微量气体交换通量。具有原位且对地表无干扰的特点, 在测定风速和风向波动的基础上得到湍流转换系数, 即“涡旋扩散系数”。微气象学方法能获得一块区域 (即通量足迹或源区) 的平均表面通量。由于测量过程基本对被测对象无影响, 因此, 微气象法比起通量箱法有许多优点, 比

如可以代表较大的空间尺度和进行长时间的测量。

测量大气微量气体通量的微气象学方法主要有3类(表3)。根据其实施的约束条件增加的顺序分别为:(1)涡旋相关法(eddy covariance method, EC);(2)弛豫涡旋积累法(relaxed eddy accumulation, REA);(3)通量梯度法(flux gradient methods),包括修正的波文比法(modified Bowen ratio, MBR)。

以上各种微气象学方法中最精确的是涡旋相关法,它是一种通过计算物理量的脉动与风速脉动的协方差获得湍流输送量(湍流通量)的方法。这种方法需要在某个特定高度上,用快速响应超声风速仪测定风速的瞬时高频波动,并且同步测定痕量气体含量高频波动,这称之为涡旋相关(一般频率为10 Hz)。当前,已经有了能够快速、简单和可靠测定一些微量气体(如CO₂和H₂O蒸汽)的设备,它们运行时无需人员照看。虽然Bauer等(2002)提议用2台激光诱导荧光光谱仪的涡旋相关法来测定汞通量,然而,现在还未见相关的报道。该提议的设备太大,并且很难运输和维护。因此,目前为止,利用涡旋相关法研究地表界面与大气间汞的交换通量难度较大。

第二种微气象学方法,弛豫涡旋积累法(REA)。当微量气体的测定只有时间分辨率较低的仪器时,在涡旋相关理论基础上发展了弛豫涡旋积累法。弛豫涡旋积累法是利用一个快速响应的垂直风速感应器(一般是超声风速计)测量垂直风速脉动,通过电磁阀系统的开合,将上升气流与下降气流以恒定的流速分别采集在2个取样袋中,再分别测定2个取样袋中气体含量。最后,通过上升气流和下降气流含量差及垂直风速的标准偏差计算通量。目前,应用REA技术在很多区域(大草原、森林林冠层、农田、填埋场、冰架、雪域等)对很多微量气体(NH₃、HNO₃、温室气体、DMS、BVOC等)和颗粒物的通量都进行了测量(Sutton et al., 2001; Zemmelink et al., 2004; Graus et al., 2006; Haapanala et al., 2006; Pattey et al., 2006; Miles et al., 2007)。但国际上,运用这种技术来研究地表界面与大气间汞的交换通量的工作也才刚刚起步。而在我国,只有北京大学朱彤教授的研究组利用这种技术测定挥发性有机污染物地表界面与大气间的交换通量(Zhu et al., 1999, 2000)。

第三种方法,通量梯度法,认为湍流以相同的条件传输不同的气体。利用这一假设,用微气象学方

法在界面以上两个或者更多的高度上同步测定汞的含量,可以用于定量垂直湍流混合速率。这些变量结合在一起就可以进行大气和界面之间的汞通量计算。作为典型的通量梯度方法,修正的波文比法(MBR)(Lindberg et al., 1995),先通过测定一种较为容易测定的示踪剂(如CO₂、H₂O)的通量、涡旋相关示踪剂以及汞在多个不同高度的浓度梯度,来计算汞的通量。这一技术依赖于精确的测定一段时间内平均的、非常小的汞含量梯度。梯度方法要求对微量气体分析检测限要比REA法高得多,因此目前该方法的精度较差。

由此可见,弛豫涡旋积累法(REA)有望成为研究地表界面与大气间汞交换通量的方法。

2000年以来,美国和加拿大科学家联合在加拿大安大略省西北试验湖区(Experimental Lake Area)的658湖,开展了纯单一汞同位素的加入实验研究(Mercury Experiment To Assess Atmospheric Loading In Canada and the United States METAALICUS项目),以确定大气汞沉降通量的变化如何影响湖泊鱼体汞含量。如图3所示,科研人员在上游高地、上游湿地和湖面分别加入3种单一汞同位素²⁰²Hg、²⁰⁰Hg和¹⁹⁸Hg,最终通过测定鱼体这3种汞同位素的过剩量来确定大气汞沉降通量的变化是通过何种渠道影响鱼体汞含量。该项目的部分成果已发表

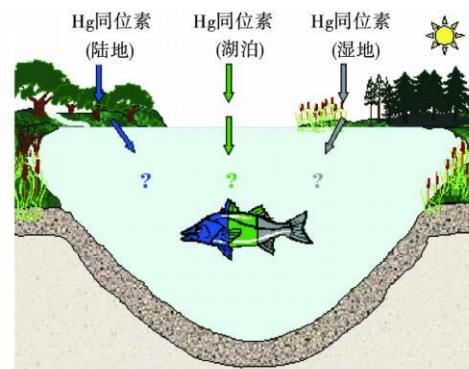


图3 METAALICUS项目利用单一汞同位素加入方法定量大气汞的直接沉降、上游高地径流及湿地径流带入湖泊的汞对鱼体汞含量的贡献

Fig. 3 In METAALICUS Project mercury isotopes were added into a lake system to quantify the contribution of Hg in fish from direct deposition, upland and wetland from the upstream of the lake.

资料来自项目网页 http://www.biogeo.uwaterloo.ca.old_site/metaalicus/metaalicus.htm

(Hinteßmann et al., 2002; Babiarz et al., 2003; Amyot et al., 2004; Branfireun et al., 2005; Chadwick et al., 2006; Poulin et al., 2006; Harris et al., 2007; Clarisse et al., 2009)。项目最终成果在美国科学院院刊 (PNAS) 上发表 (Harris et al., 2007)。研究证实, 大气汞沉降通量的变化会改变鱼体汞的含量。同时, 如前所述, Erickson 等 (2005) 利用单一汞同位素加入法, 研究了大气湿沉降的汞对沙漠地区土壤汞的再释放通量的影响。这种单一汞同位素的加入技术为研究植被与大气汞的交换机理及大气汞沉降再释放的控制因素提供了很好的技术参考。随着技术手段的进步, 目前已具备了开展地表自然排汞及先前排汞沉降后的再释放过程、机理和通量研究的各方面条件。

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