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新疆黄山—镜儿泉铜镍成矿带岩浆作用与区域走滑构造的关系

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摘要:中国产于造山带的铜镍硫化物矿床占有非常重要的地位,例如中亚造山带南缘的一系列矿床及近年来东昆仑造山带发现的夏日哈木超大型镍钴矿床,其探明的镍金属总储量超过 300×10^4 t, 约占中国镍金属总储量的 $1/3$,也是世界上造山带发现这类矿床最多的国家。为什么在造山带也能够形成大型—超大型铜镍硫化物矿床?与其相关的大规模幔源岩浆作用有什么特点,发生在造山带构造演化的哪个阶段?是什么机制导致了这样剧烈的镁铁质岩浆作用?对这些关键问题的认识还很模糊。以大中型铜镍硫化物矿床分布最为集中的新疆北天山黄山—镜儿泉成矿带为例,在对国内外研究成果综合归纳和分析的基础上,对上述问题进行探讨。结果表明:准噶尔—哈萨克斯坦地块与塔里木地块碰撞过程中伴随区域性右行走滑,加剧了俯冲洋壳断离和软流圈上涌,并为源自软流圈及交代地幔部分熔融的镁铁质岩浆上侵提供了良好的通道,这些作用的相互叠加为黄山—镜儿泉成矿带的形成创造了很好的条件。

关键词:岩浆铜镍硫化物矿床;幔源岩浆作用;造山带;区域走滑构造;软流圈上涌;地幔交代作用;板块碰撞;俯冲洋壳断离

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Magmatism of Huangshan-Jing'erquan Ni-Cu Ore Deposit Belt and Relationship with Regional Strike-slip Structure in Xinjiang, China

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Abstract: The Ni-Cu sulfide deposits in orogenic belts are very important in China. Nickel metal reserves of the Ni-Cu sulfide deposits discovered in orogenic belts, China, including a series of deposits in the southern margin of Central Asian orogenic belt and the large Xiarihami Ni-Co deposit in East Kunlun orogenic belt, are more than 300×10^4 t. These discoveries make China having the most abundant magmatic sulfide deposits in orogenic belts over the world. However, some important issues on the formation of such deposits have not been well understood. For

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instance, what results in the formation of those large and super-large Ni-Cu sulfide deposits in the orogenic belts? What are the features of the related magmatism? When the magmatism occurs during tectonic evolution of the orogenic belts? What is the mechanism leading the extensive mafic magmatism? The above issues predominantly based on the recent studies of Huangshan-Jing'erquan Ni-Cu sulfide deposit belt as well as Permian regional strike-slip structure in northern Tianshan, Xinjiang were focused on. The results show that the collision between Junggar-Kazakhstan block and Tarim block along with regional dextral strike-slip structure in Permian results in break-off of the subduction oceanic slab and upwelling of the asthenosphere; the upwelling of the asthenosphere causes extensive melting of itself, and the metasomatized wedge mantle produces mafic magma; the strike-slip faults provide the pass way for the ascending mafic magmas and create spaces for the magmatic sulfide mineralized intrusions in the crust of Huangshan-Jing'erquan Ni-Cu sulfide deposit belt.

Key words: magmatic Ni-Cu sulfide deposit; mantle-derived magmatism; orogenic belt; regional strike-slip structure; upwelling of the asthenosphere; mantle metasomatism; plate collision; break-off of subduction oceanic slab

0 引言

世界上已发现的超大型岩浆硫化物矿床大多产于大火成岩省或大陆裂谷中^[1-7],但在北美洲、欧洲、澳大利亚、非洲及亚洲的若干造山带也陆续发现了大中型岩浆硫化物矿床^[8-13],特别是中国新疆北天山黄山—镜儿泉成矿带一系列大型铜镍硫化物矿床的发现,以及东昆仑造山带新近发现的夏日哈木超大型镍钴矿床(镍金属储量超过 100×10^4 t)^[14-21],意味着除大火成岩省和大陆裂谷以外,造山带也是有利于铜镍硫化物成矿的构造背景。然而,由于造山带铜镍硫化物矿床的金属品位一般较低,单个矿床的规模也较小,所以其找矿工作并没有受到应有的重视。考虑到全球造山带的面积远远大于大火成岩省的面积,造山带仍然有巨大铜镍硫化物矿床的找矿潜力,但是,这些矿床具有什么样的特殊成矿规律尚需进一步研究。

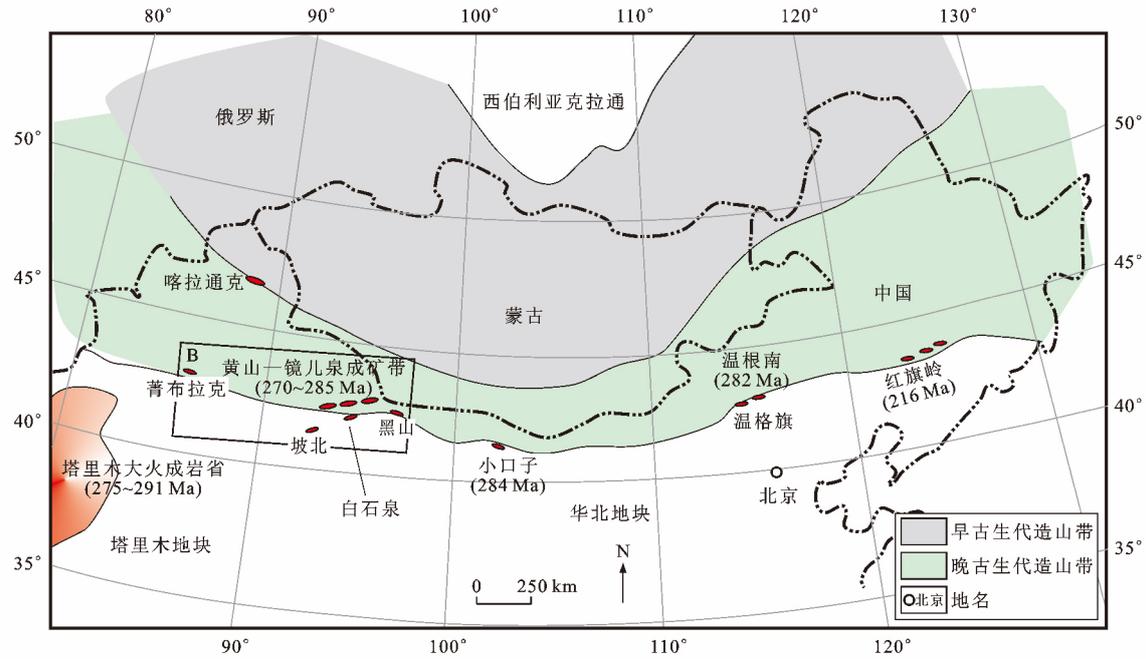
中亚造山带是世界上罕见的、由多条重要成矿带构成的多金属成矿域,包括中国阿尔泰铜-金-多金属矿带、蒙古南部斑岩铜矿带、北哈萨克斯坦金-铀成矿带、中哈萨克斯坦铁-锰-铜-多金属和稀有金属成矿带。这些成矿带分布着许多世界知名的大型、超大型矿床,例如蒙古南部世界最大的在采斑岩铜矿床——欧玉陶勒盖矿床。

中国北部从新疆北部经内蒙古至东北地区是中亚造山带的重要组成部分,也被称为兴蒙造山带。与周围其他国家比较,岩浆硫化物矿床是该区域特有的优势矿种之一,如东段的吉林红旗岭

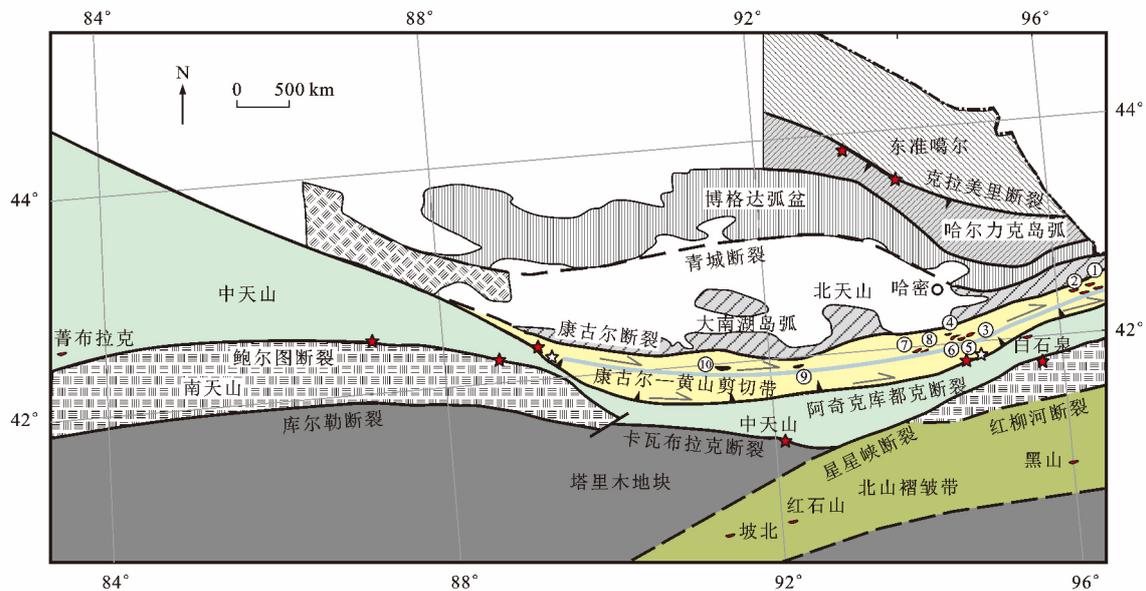
大型铜镍矿床及中段的内蒙古温格旗和尔布图等矿化岩体,而更多的大中型铜镍硫化物矿床发现于新疆北部及相邻的甘肃西部。新疆北部已探明的镍金属储量超过 200×10^4 t,是中国仅次于金川的第二大镍资源基地,更是世界上规模最大的造山带岩浆铜镍硫化物矿集区。其中黄山—镜儿泉成矿带又是新疆北部大型铜镍硫化物矿床最集中、空间分布最有规律的铜镍硫化物成矿带。然而,为什么数千千米的中亚造山带南缘在新疆北部岩浆硫化物成矿作用尤其强烈和集中?是什么条件、过程和机制使得黄山—镜儿泉成矿带上形成了时间上比较集中、空间上比较密集且呈带状分布的若干大中型矿床?对这些问题进行系统和深入的探讨对于认识中亚造山带成矿规律,特别是造山带铜镍硫化物矿床的形成规律具有重要意义。近年来,不少学者借助于岩石和矿床地球化学手段对黄山—镜儿泉成矿带多个矿床成因开展了大量研究^[22-34],但对岩浆作用的区域构造背景认识还较为薄弱。本文试图在上述研究成果的基础上,对黄山—镜儿泉成矿带中导致铜镍硫化物成矿的镁铁质岩浆活动与俯冲-碰撞过程及大规模区域走滑构造的潜在联系进行探讨。

1 区域地质概况

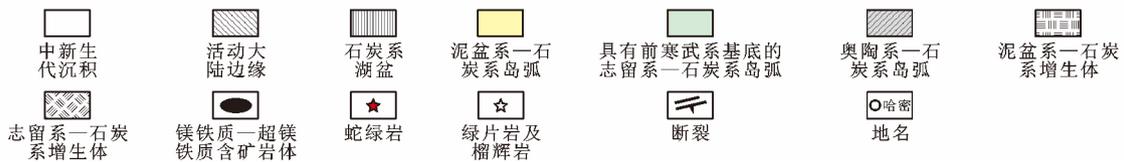
新疆北部处于中亚造山带南部,由北向南可以划分为阿尔泰地块、准噶尔地块、天山褶皱带和塔里木地块等构造单元,天山褶皱带又进一步分为北天山、中天山和南天山(图1)。主要的铜镍硫化物含



(a) 中亚造山带地质简图



(b) 新疆北部构造单元和主要岩浆硫化物矿床分布



图(a)括号中是镁铁质—超镁铁质含矿岩体的锆石 U-Pb 年龄;①为图拉尔根;②为葫芦;③为黄山东;④为香山;⑤为黄山南;⑥为黄山;⑦为土墩;⑧为二红洼;⑨为海豹滩;⑩为恰特卡尔塔格;图件引自文献[30]、[31]、[35]~[37],有所修改;主要镁铁质—超镁铁质含矿岩体的锆石 U-Pb 年龄见表 1

图 1 中亚造山带地质简图及新疆北部构造单元和主要岩浆硫化物矿床分布

Fig. 1 Simplified Sketch Map of Central Asian Orogenic Belt and Distribution of Tectonic Units and Major Magmatic Sulfide Deposits in Northern Xinjiang

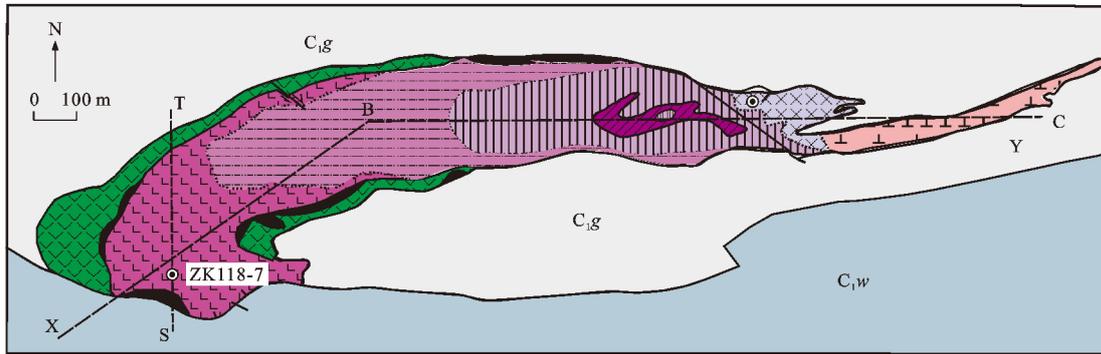
矿岩体分布在准噶尔地块北缘(喀拉通克)、北天山东部的天宇和白石泉)以及塔里木地块北缘的北山褶皱带(西部的坡北、红石山和罗东,东部的黑山),

这些含矿岩体在各构造单元中的分布都大致平行于区域性大断裂或板块缝合线(图1)。近年来较为精确的锆石 U-Pb 年龄表明除个别含矿岩体形成于晚泥盆世(黑山)和志留世(菁布拉克)外,多数含矿岩体形成于早二叠世(270~290 Ma)。黄山—镜儿泉成矿带含矿岩体的含矿岩相多形成于早二叠世(280~285 Ma,除图拉尔根和黄山东外),而非含矿岩体及岩体的不含矿岩相(如黄山闪长岩)形成较晚(270~280 Ma)(表1)。这种现象与俄罗斯西伯利亚大火成岩省 Noril'sk 等含矿岩体以及中国峨眉大火成岩省岩浆硫化物含矿岩体有明显区别,这些由地幔柱岩浆活动形成的含矿岩体不同岩相具有一致的形成年龄及显著的分异演化关系。

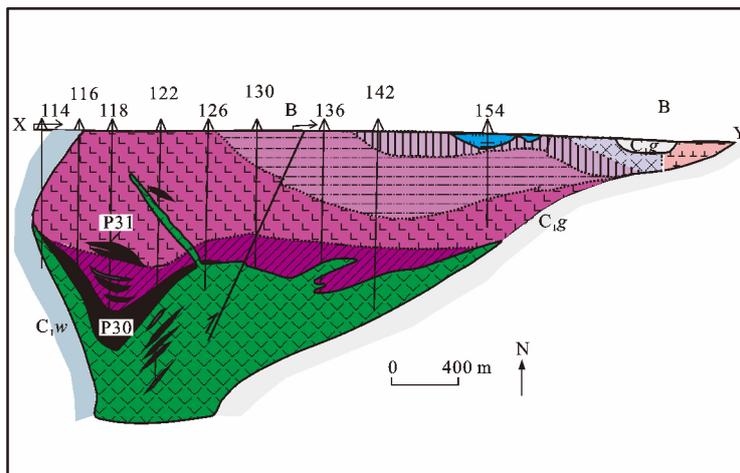
黄山—镜儿泉成矿带长约 450 km、宽约 50 km,位于区域性的阿奇克库都克—沙泉子断裂和康古尔断裂之间,岩体平行于这两个断裂分布(图1)。

该成矿带含矿岩体具有以下特点:①均为镁铁质—超镁铁质岩体,含原生普通角闪石及黑云母;②具有岩浆通道成矿和多次岩浆侵入的特点,岩相间的交切关系显示含矿岩相形成较早;③矿体主要赋存于超镁铁质岩相中;④Ni 平均品位一般低于 0.6%,个别高于 1%,铂族(PGE)平均品位低于 0.1×10^{-6} ,个别为 $(0.5 \sim 1.0) \times 10^{-6}$;⑤超镁铁质岩相中橄榄石 Fo 牌号较高(80~86),岩石大离子亲石元素相对富集,高场强元素相对亏损,具有 Nb、Ta 负异常和 Pb 正异常的特点,显示与成矿有关的高镁玄武质岩浆与交代地幔的部分熔融有关;⑥地壳物质同化混染对成矿具有重要意义^[22-48]。

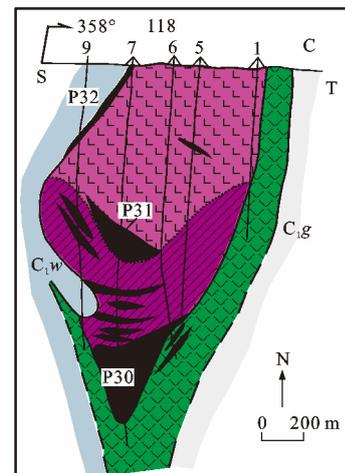
根据各岩相的空间分布、接触和交切关系,可以分辨出黄山岩体主要有 3 次岩浆侵入(图2):第一次岩浆侵入形成了岩体东部细粒二辉橄榄岩上悬体,代表了岩浆分异系列的底部,是后期剥蚀的残留



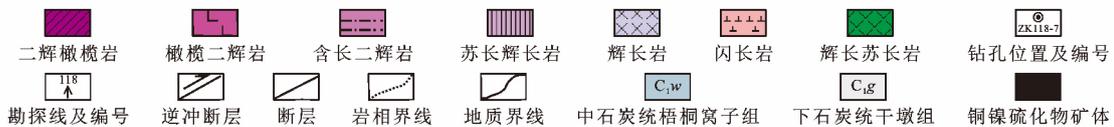
(a) 平面地质简图



(b) 横剖面地质图



(c) 主要勘探线剖面



图件引自文献[38],有所修改;P30、P31、P32 为矿体编号

图2 黄山岩体平面地质简图、横剖面地质图和主要勘探线剖面

Fig. 2 Simplified Geological Map, Transverse Section Geological Map and Main Cross Section of Huangshan Intrusion

表 1 中亚造山带南缘主要镁铁质—超镁铁质含矿岩体锆石 U-Pb 年龄及铜镍金属储量

Tab. 1 Zircon U-Pb Ages and Ni-Cu Metal Reserves of the Main Mafic-ultramafic Ore-bearing Intrusions in the Southern Margin of Central Asian Orogenic Belt

构造单元	岩体	U-Pb 年龄测试方法	年龄/Ma	镍金属储量 (10 ⁴ t)或规模	铜金属储量 (10 ⁴ t)或规模	文献来源
准噶尔 地块北缘	喀拉通克	苏长岩锆石 SHRIMP 法	287.0±5.0	25.00(大型)	42.00(大型)	[49]
北天山 黄山— 镜儿泉 成矿带 (由东向 西排列)	图拉尔根	辉长岩锆石 SHRIMP 法	300.0±3.2	11.00(大型)	6.00(大型)	[50]
	圪塔山口	辉长岩锆石 SIMS 法	283.0±1.9	小型	小型	[51]
	葫芦	辉长闪长岩锆石 LA-ICP-MS 法	282.0±1.2	8.00	4.00	[52]
	黄山东	橄榄苏长岩锆石 SHRIMP 法	274.0±3.0	36.00(大型)	17.00(大型)	[49]
	黄山南	辉长岩锆石 SHRIMP 法	283.0±1.4	新增储量估计为大型		[53]
	黄山	闪长岩锆石 SHRIMP 法(弱矿化, 与辉长(苏长)岩呈侵入接触关系)	269.0±2.0	32.00(大型)	20.00(大型)	[22]
		辉长(苏长)岩锆石 LA-ICP-MS 法	284.0±2.5			[54]
	香山	苏长辉长岩锆石 SHRIMP 法	285.0±1.2、 280.0±1.1	4.00	2.00	[32]
	二红洼	橄榄辉长岩锆石 LA-ICP-MS 法	283.0±1.5	弱矿化	弱矿化	[46]
	白鑫滩	辉橄岩及橄榄辉长岩锆石 LA- ICP-MS 法	278.0±2.6、 287.0±3.0	弱矿化	弱矿化	[55]及邓宇峰未 发表数据
海豹滩	辉长岩锆石 SHRIMP 法	269.0±3.2	未见矿化	未见矿化	[56]	
恰特卡尔塔格	辉长岩锆石 SHRIMP 法	277.0±1.6	未见矿化	未见矿化	[56]	
中天山	白石泉	辉长岩锆石 SHRIMP 法	284.0±8.0	9.40	7.00	[26]、[57]
		辉长岩锆石 LA-ICP-MS 法	281.0±0.9			
	天宇	辉长岩锆石 LA-ICP-MS 法	290.0±3.4			[58]
	菁布拉克	辉石闪长岩锆石 SHRIMP 法	431.0±6.0	0.42	0.11	[44]
北山褶 皱带	坡北	辉长岩锆石 SHRIMP 法	278.0±2.0、 289.0±10.0	50.0(128.0) (超大型)	0.4(0.2) (超大型)	[26]、[59]
	红石山	橄榄辉长岩锆石 SHRIMP 法	282.0±2.6			[60]
	黑山	辉长岩锆石 SHRIMP 法	357.0±4.0	12(大型)	4(大型)	[47]
内蒙古	温格旗	辉石岩锆石 SHRIMP 法	399.0±4.0			[61]
	尔布图	橄榄单辉岩锆石 SHRIMP 法	294.2±2.7			[62]

部分;第二次岩浆侵入形成了黄山岩体的主体,自下而上由二辉橄榄岩、橄榄二辉岩、含长石二辉岩、苏长辉长岩、辉长岩和闪长岩组成,它们之间呈渐变过渡关系,构成了比较完整的分异系列;第三次岩浆侵入形成了岩体底部的辉长苏长岩。矿化主要产于第二次岩浆侵入形成的二辉橄榄岩底部。近年来获得的锆石 U-Pb 年龄为 (269±2) Ma^[22] 和 (284.0±2.5) Ma^[54], 这些锆石 U-Pb 年龄的差异很可能代表了第三次和第二次岩浆侵入的时间,这说明该成矿带镁铁质岩浆活动的时间比较长,有异于地幔柱幔源岩浆活动时间非常集中的特点。岩相分

布表明多次侵入的岩浆总是沿着基本一致的岩浆通道侵入,新的岩浆一般沿前一次岩浆形成的岩相底部侵入,类似现象也发生在黄山东和土墩等含矿岩体,这也是黄山—镜儿泉成矿带含矿岩体的突出特点之一。

2 成矿带与俯冲-碰撞过程的关系

表 1 中锆石 U-Pb 年龄数据表明,黄山—镜儿泉成矿带的铜镍硫化物矿床成矿具有“爆发式成矿”的特点。对具有这一特征的矿床成矿有两类认识:与俯冲-碰撞过程有关;与塔里木地幔柱有关。

中亚造山带南缘古亚洲洋最终闭合时间究竟是晚石炭世还是二叠纪或三叠纪尚有争议^[63-67],因此,对黄山—镜儿泉成矿带铜镍硫化物矿床究竟形成于俯冲阶段、碰撞阶段还是碰撞后阶段仍然存在较大分歧。Zhang等认为这些含矿岩体形成于板块俯冲阶段^[28-31,46,49,65-75]。

2.1 与塔里木地幔柱的成因联系

Pirajno等根据新疆北部若干铜镍硫化物含矿岩体年龄与塔里木大火成岩省侵入岩年龄的相似性,认为塔里木地幔柱为这些矿床的形成提供了热源^[76]。这种观点符合地幔柱可能为地幔高度部分熔融提供热源的基本原理。但是,该观点在以下4个方面还值得商榷:①塔里木玄武岩分选出的锆石 U-Pb 年龄((289.5±2.0)Ma^[77]、(288.0±2.0)Ma^[77]、(291.0±4.0)Ma^[78])说明玄武岩活动主要发生在290 Ma左右,延续时间大约5.5 Ma,略早于新疆早二叠世铜镍硫化物含矿岩体的锆石 U-Pb 年龄(表1),而中基性岩体(275~280 Ma)的形成略晚于铜镍硫化物含矿岩体;②新疆铜镍硫化物含矿岩体发育的几个构造单元不仅远离塔里木大火成岩省,而且这些构造单元发现的早二叠世玄武岩地球化学性质也与塔里木大火成岩省镁铁质侵入岩和玄武岩相去甚远^[79-81];③Zhang等根据岩石地球化学特征的差异,认为喀拉通克和黄山岩体与塔里木地幔柱之间没有任何成因联系^[81-82],最接近塔里木地幔柱的坡北、白石泉和天宇岩体与塔里木大火成岩省镁铁质—超镁铁质岩体及岩墙的地球化学特征也有非常明显的区别^[30];④更重要的是,远离塔里木大火成岩省的地区也发现了280~290 Ma的镁铁质—超镁铁质含矿岩体(如约2000 km以外华北地块北缘的内蒙古温根南中型铜镍硫化物矿床(锆石 U-Pb 年龄为282 Ma)被认为与梭楞克尔俯冲带向南俯冲有关^[83];甘肃阿拉善地块含铜镍矿化的小口子辉长岩体的锆石 U-Pb 年龄为(284.4±3.5)Ma(焦建刚未发表数据)),这些岩体的微量元素组成也与新疆北部铜镍硫化物含矿岩体一致,如果这些岩体与塔里木地幔柱有成因联系,则塔里木大火成岩省的面积将远大于俄罗斯西伯利亚大火成岩省,但塔里木玄武岩的最大厚度刚超过500 m,有多个沉积夹层,且向塔里木盆地中部逐渐尖灭^[84-85],远远小于西伯利亚玄武岩数千米的厚度^[86]。

2.2 俯冲洋壳断离与突发性带状镁铁质岩浆活动

板块聚合边界的构造-岩浆-成矿作用研究一直是地质学最为活跃的研究领域之一,近年来许多重

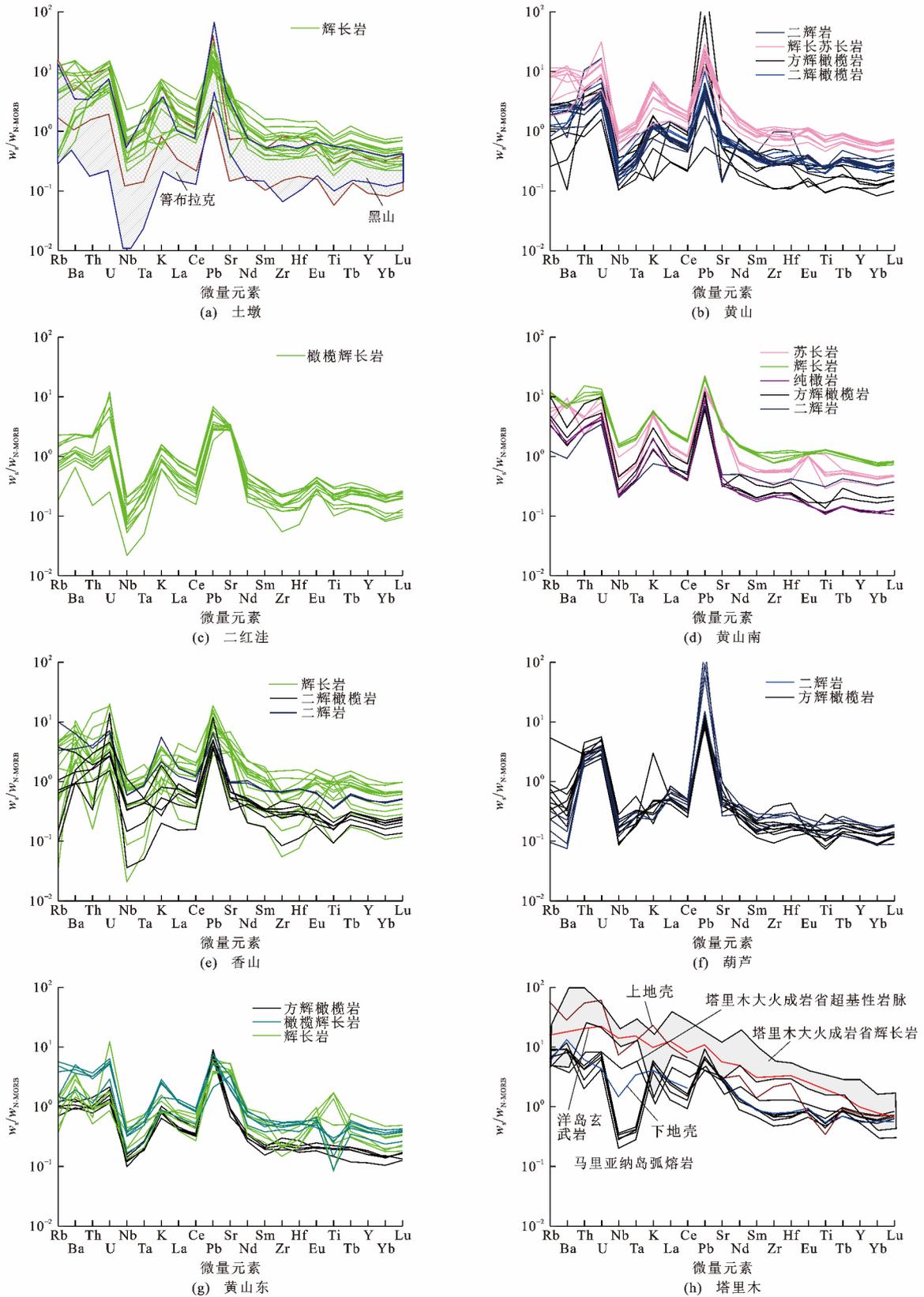
要进展正在不断加深对俯冲-碰撞过程中岩浆活动的认识。尽管岛弧和活动大陆边缘岩浆活动以钙碱性的中酸性岩浆为主,但俯冲板片后撤、板片撕裂或洋脊俯冲都可能导致软流圈上涌,并产生拉斑玄武质岩浆活动^[87-90]。特别是碰撞和后碰撞阶段,俯冲板片失速断离、岩石圈地幔拆沉、软流圈热对流对岩石圈地幔侵蚀减薄都可能导致软流圈上涌及镁铁质岩浆活动,其中第一种机制可能导致突发性带状镁铁质岩浆活动,而后两种机制会导致面状幔源岩浆活动^[91-92]。墨西哥近1000 km的新近纪玄武质火山岩带年龄的规律变化表明俯冲板片侧向撕裂的速度为100~250 km·Ma⁻¹^[90],说明板片撕裂断离也可以导致“爆发式”线状幔源玄武质岩浆活动。

2.3 含矿岩体地球化学特征的启示

近年来,Song等研究认为黄山—镜儿泉成矿带上的含矿岩体全岩微量元素组成普遍存在 Nb、Ta 负异常(图3)^[30-31]。火山岩 Nb、Ta 负异常可能是幔源岩浆遭受地壳物质同化混染的结果,也可能说明受俯冲事件影响交代地幔源区来源的岩浆。然而,对于镁铁质—超镁铁质侵入岩而言,全岩 Nb、Ta 负异常可能与堆积相中存在亏损 Nb、Ta 的矿物有关。因此,仅用全岩正常型洋中脊玄武岩标准化微量元素蛛网图判断镁铁质—超镁铁质侵入岩地幔源区特征有一定的不确定性,而微量元素含量比值能较好地反映岩浆的地球化学特征。进一步的全岩微量元素含量比值图解(图4)说明黄山—镜儿泉成矿带含矿岩体的微量元素组成不能用地壳物质同化混染解释,代表了一种受俯冲事件改造的地幔源区特征。

黄山、黄山东及黄山南岩体纯橄岩和方辉橄岩中包裹在高镁橄榄石牌号(Fo)橄榄石内的铬尖晶石成分表明软流圈地幔也强烈参与了部分熔融(图5)。需要说明的是,黄山东岩体方辉橄岩中包裹铬尖晶石的橄榄石 Fo 牌号较低(81~86),说明其较高的 TiO₂ 含量和较低的 Al₂O₃ 含量与较强烈的分离结晶有关,但并不能说明其具有与洋岛玄武岩相同的地幔源区^[94]。

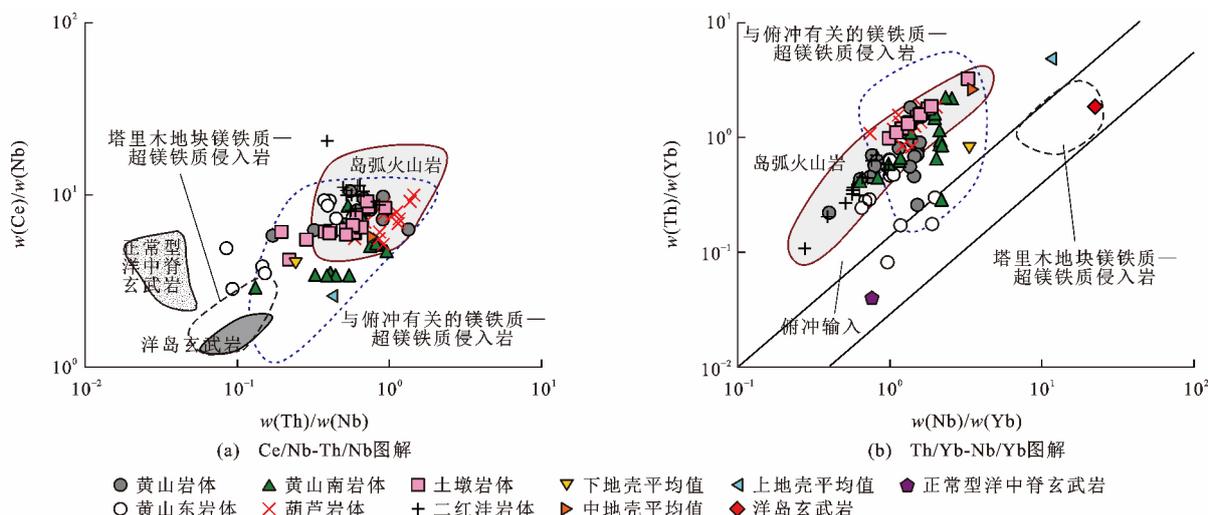
Song等提出原始岩浆的形成与同碰撞阶段俯冲板片断离导致的软流圈上涌密切相关^[31]。受岩石密度影响,同碰撞阶段密度较小的陆壳向下俯冲的深度十分有限,必然导致榴辉岩化的俯冲洋壳与密度较小的俯冲陆壳发生断离。当这种断离深度较浅时,上涌的软流圈本身就可能发生减压部分熔融,



图件引自文献[31];塔里木大火成岩省超基性岩脉及辉长岩数据引自文献[79];菁布拉克和黑山岩体数据引自文献[44]和[47];
 w_p 为样品含量; w_{N-MORB} 为正常型洋中脊玄武岩含量; 同一图中相同线条对应不同样品

图 3 主要含矿岩体正常型洋中脊玄武岩标准化微量元素蛛网图

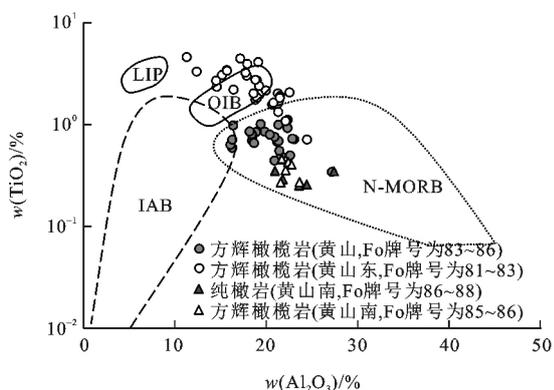
Fig. 3 N-MORB-normalized Trace Element Spider Diagrams of the Dominant Ore-bearing Intrusions



图件引自文献[31]; 洋岛玄武岩、洋中脊玄武岩、岛弧火山岩及与俯冲有关的镁铁质—超镁铁质侵入岩数据引自文献[9]和[93]; 塔里木地块镁铁质—超镁铁质侵入岩引自文献[79]和[80]

图 4 主要含矿岩体全岩 Ce/Nb-Th/Nb 图解和 Th/Yb-Nb/Yb 图解

Fig. 4 Diagrams of Ce/Nb-Th/Nb and Th/Yb-Nb/Yb of the Dominant Ore-bearing Intrusions



图件引自文献[31]; 底图引自文献[94]; N-MORB 为正常型洋中脊玄武岩; LIP 为层状侵入体橄辉岩; OIB 为洋岛玄武岩; IAB 为岛弧玄武岩

图 5 包裹在高镁橄榄石牌号 (Fo) 橄榄石内的铬尖晶石 TiO₂-Al₂O₃ 投影图

Fig. 5 Plot of TiO₂ Against Al₂O₃ of the Chromite-hosted in the Olivines with High Fo

同时导致上覆交代地幔的加热和部分熔融,两种岩浆的混合形成了黄山—镜儿泉成矿带的原始岩浆。这个观点能够较好地解释黄山—镜儿泉成矿带含矿岩体与形成于俯冲阶段的西天山志留世菁布拉克和北山褶皱带晚泥盆世黑山岩体具有非常相似的岩石组合、矿物组成和地球化学特征(表 1, 图 1、3)。Yang 等对菁布拉克和黑山岩体铜镍硫化物矿床的研究表明,俯冲板片撕裂导致的软流圈上涌也可以形成与岩浆硫化物成矿有关的玄武质岩浆,并不需要地幔柱提供热源^[44,47-48]。Li 等在东昆仑造山带夏日哈木超大型镍钴矿床的研究也证明地幔柱背景

并非大规模岩浆硫化物成矿作用的先决条件^[18,21]。

3 成矿带与天山构造带早二叠世大规模走滑的关系

20 世纪 80 年代的地球物理研究表明,康古尔断裂南、北两侧存在急剧的重力及地磁梯度差异,反映出地壳结构和岩性密度的差异,这与其两侧古生代地层和岩性组合的差异是一致的^[38]。这些区域地质特征表明康古尔断裂为区域性深断裂,黄山—镜儿泉镁铁质—超镁铁质岩带产于该断裂南侧。李德惠等的构造地质学研究已经在黄山—镜儿泉成矿带识别出韧性剪切构造^[38],包括拉伸线理、沉积岩砾石的塑性变形等,但由于当时缺乏有效的构造年代学分析方法,所以对韧性剪切构造发生的时间认识不够清晰。

杨兴科等将东天山构造带北部的巨型韧性剪切带称为秋格明塔什—黄山巨型韧性剪切带,其主要标志是石炭系火山岩和沉积岩不仅发生了绿片岩相变质,还普遍发生了脆—韧性变形,如糜棱岩化、拉伸线理、剪切褶皱、沉积砾石塑性拉伸、变形石香肠、镜下矿物波状消光等^[95]。该韧性剪切带是准噶尔—哈萨克斯坦地块与塔里木地块之间的俯冲—碰撞产物,右行剪切位移约为 75 km。

Wang 等根据古地磁研究认为在晚石炭世—早二叠世时期,伊利—准噶尔地块相对于塔里木地块发生了上千千米的右行走滑,形成了巨型的韧性剪切带,并将该巨型剪切带的东段称为康古尔—黄山

韧性剪切带^[96]。Wang等发现该构造带上340 Ma左右的石炭纪花岗岩没有韧性剪切变形,而二叠纪花岗岩遭受了显著的韧性剪切变形,显示同构造侵位的特点,因此,认为二叠纪侵入岩的分布可能受韧性剪切构造的控制^[37]。陈文等根据韧性剪切带中岩石全岩、钾长石、斜长石的Ar-Ar年代学研究认为该韧性剪切带早期(280~300 Ma)以推覆剪切为主,可能与板块俯冲-碰撞有关,晚期(242~262 Ma)以右行走滑剪切为主,反映了碰撞后陆内走滑、抬升过程^[97]。

根据韧性剪切构造的分布范围和地质特点,特别是黄山和黄山东岩体边缘及内部变形特征,Branquet等对黄山和黄山东岩体边缘及围岩的韧性剪切构造及岩体内部的断裂构造进行了细致分析,认为是区域性剪切作用形成的次一级张性构造为幔源岩浆上升提供了通道,并为岩浆房的形成提供了空间^[36]。黄山和黄山东岩体经历的多阶段岩浆侵入与区域性剪切作用是同时发生的,形成了黄山岩体的“蝌蚪状”和黄山东岩体的“菱形”形态,以及它们向下收敛的楔形纵剖面^[36]。这些特点符合岩浆通道系统成矿的特点,但都有别于大陆裂谷环境形成的镁铁质—超镁铁质侵入岩的地质特征^[98-99]。这些构造地质学研究成果为进一步理解黄山—镜儿泉成矿带的幔源岩浆作用诱发机制提供了新的依据。

4 成矿带幔源岩浆作用的可能机制

黄山—镜儿泉成矿带与成矿有关的玄武质岩浆作用是早二叠世康古尔—雅满苏岛弧与中天山地块碰撞导致俯冲板片断离、软流圈上涌及地幔部分熔融的结果,区域走滑构造则为玄武质岩浆上侵提供了通道,为含矿岩体的形成创造了空间。

同碰撞阶段发生大规模的区域性走滑剪切时,不仅使上述俯冲洋壳的断离更加容易,也可能有助于软流圈地幔及交代地幔的减压熔融,剪切走滑产生的局部拉伸可能为这种幔源岩浆的上升提供了非常好的通道,也在地壳为含矿岩体的形成创造了良好空间,从而在近500 km狭长的黄山—镜儿泉成矿带上形成若干含矿岩体。黄山和黄山东岩体独特的形态说明岩体就位与走滑剪切空间的确存在密切的成因联系^[36-37]。

黄山—镜儿泉成矿带的这些地质和地球化学特征以及区域构造背景说明,在聚合板块边缘,一些特殊的地质过程也可以产生有利于大规模地幔部分熔

融和岩浆硫化物成矿的条件,这为在其他造山带寻找大型—超大型岩浆硫化物矿床提供了科学依据。需要指出的是,不同聚合板块边缘的地质作用过程往往存在区别,要根据具体情况进行细致的分析和研究,不能一概而论。

5 结 语

(1)新疆黄山—镜儿泉铜镍硫化物成矿带是一个长约450 km、宽约50 km的镁铁质—超镁铁质岩带,岩浆作用具有“爆发式”特点。

(2)最新的区域构造研究表明,黄山—镜儿泉成矿带产于区域性右行走滑构造带上,该右行走滑构造的活动时间为晚石炭世—早二叠世,与准噶尔—哈萨克斯坦地块和塔里木地块碰撞时间非常吻合。

(3)准噶尔—哈萨克斯坦地块与塔里木地块碰撞以及区域性走滑构造的叠加很可能是导致俯冲洋壳断离、软流圈上涌和上地幔部分熔融的关键诱因,从而在碰撞带发生了“爆发式”的幔源岩浆作用,为铜镍硫化物成矿提供了物质基础。

(4)区域走滑形成的局部拉张为幔源岩浆上升和含矿岩体的形成提供了通道和空间,为黄山—镜儿泉成矿带的形成提供了较为特殊的地质背景。

(5)造山带发生及发展过程中,俯冲-碰撞也是岩浆硫化物大规模成矿的有利背景。

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